

Annual Cloud Seeding Report
High Uintas Program
2021-2022 Winter Season

Prepared For:

Duchesne County Water Conservancy District
Uintah Water Conservancy District
State of Wyoming
Metropolitan Water District of Southern California
State of Utah, Division of Water Resources

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EXECUTIVE SUMMARY

The Duchesne County Water Conservancy District and the Uintah Water Conservancy District were joined by the Central Utah Water Conservancy District in supporting an operational seeding project, beginning in the 2002-2003 winter season. The intended target area of this program has been the southern slope of the Uinta Mountains above 8,000 feet. The High Uintas program currently utilizes 20 ground-based, manually operated Cloud Nuclei Generator (CNG) sites, burning a 2% Silver Iodide solution. Some sites established for the adjacent Western Uintas seeding program are also utilized to target the High Uintas. In addition, a remotely operated site located at Moon Lake has been added to the program. The goal of the seeding program is to augment wintertime snowpack/precipitation over the seeded watersheds. Cost sharing for the seeding program is provided by the Utah Division of Water Resources, with additional funds from the Lower Colorado River Basin States providing for an early-season extension to the seeding program.

Precipitation and snowfall were again mostly below normal during the 2021-2022 winter season. However, the Uintas region fared pretty well this season compared to many surrounding areas. As of April 1, 2022, SNOTEL observations for the Natural Resource Conservation Service showed snow water content averaging about 88% of normal (median) for the Duchesne Basin and about 91% of normal for sites in the portions of the Uinta Range that compose the Green River Basin. Water year precipitation percentages were 107% of the median for the Duchesne Basin and around 105% of normal for sites in the Green River Basin. These above average (median) values for water year precipitation were due in part to a very wet month of October, prior to any significant snowpack accumulation.

A total of 1,617.5 CNG hours were conducted during 23 storm periods for the core High Uintas program this season, out of a maximum budgeted 2,750 hours. An additional 115 hours of seeding were conducted (during 2 storm periods) in November for the Lower Basin States sponsored extension period. There were no seeding suspensions during the 2021-2022 season.

Evaluations of the effectiveness of the cloud seeding program were made for both the past winter season and for all seeded seasons combined. These evaluations utilize SNOTEL records collected by the Natural Resources Conservation Service (NRCS) at selected sites within and surrounding the seeded target area, as well as some seasonal streamflow data. Analyses of the effects of seeding on target area precipitation, snow water content, and streamflow have been conducted for this seeding program, utilizing target/control comparison techniques.

As summarized in Section 5 of the report, determination of the exact seeding effects in the High Uintas is particularly challenging for a variety of reasons. **Based on a review of nearly 2-decades worth of data, NAWC has estimated that the seeding program is generating approximately a 3-5% seasonal increase in seasonal snowpack for this program. This equates to an approximate program yield of 36,000 additional acre-feet of annual runoff as well as significant increases to ground water/aquifer recharge.**

WEATHER MODIFICATION OVERVIEW

The Science

The cloud-seeding process aids precipitation formation by enhancing ice crystal production in clouds. When the ice crystals grow sufficiently, they become snowflakes and fall to the ground.

Silver iodide has been selected for its environmental safety and superior efficiency in producing ice in clouds. Silver iodide adds microscopic particles with a structural similarity to natural ice crystals. Ground-based and aircraft-borne technologies can be used to add the particles to the clouds.

Safety

Research has clearly documented that cloud seeding with silver-iodide aerosols shows no environmentally harmful effect. Iodine is a component of many necessary amino acids. Silver is both quite inert and naturally occurring, the amounts released are far less than background silver already present in unseeded areas.

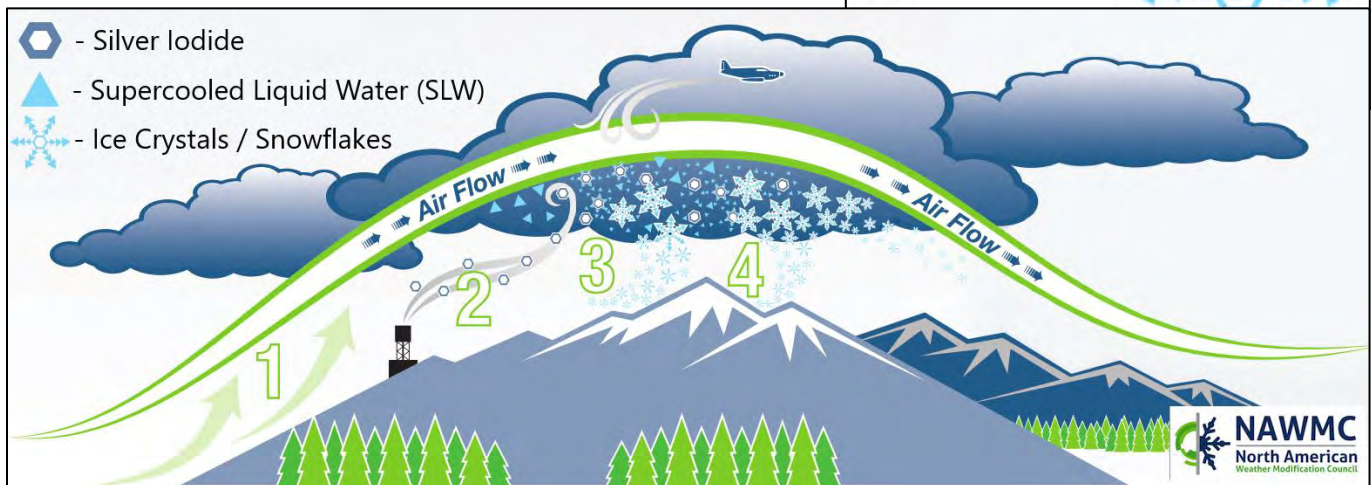
Effectiveness

Numerous studies performed by universities, professional research organizations, private utility companies and weather modification providers have conclusively demonstrated the ability for Silver Iodide to augment precipitation under the proper atmospheric conditions.

Silver-iodide crystals have a shape similar to ice crystals and provide a “seed” or nucleus for ice formation when placed in a cloud.

Droplets of supercooled water in the cloud attach to the silver iodide and form ice crystals.

Ice crystals grow until they acquire enough mass, form a snowflake, and fall toward Earth.



STATE OF THE CLIMATE

As reported last year, every ten years, the National Oceanic and Atmospheric Association (NOAA) releases a summary of various U.S. weather conditions for the past three decades to determine average values for a variety of conditions, including, temperature and precipitation. This is known as the U.S. Climate normal, with a 30-year average, representing the “new normal” for our climate. These 30-year normal values can help to determine a departure from historic norms and identify current weather trends.

The recently released 30-year average ranges from 1990 – 2020. Images in Figure 1 and 2 show how each 30-year average for the past 120 years compares to the composite 20th century average for temperature and precipitation.

For the western U.S., the 1990-2020 average shows much warmer than average temperatures. When comparing precipitation for the past 30 years to both the previous 30-year average and the 1901-2000 average, the American Southwest (including portions of Utah, Arizona, California and Nevada) has seen as much as a 10% decrease in average annual precipitation.

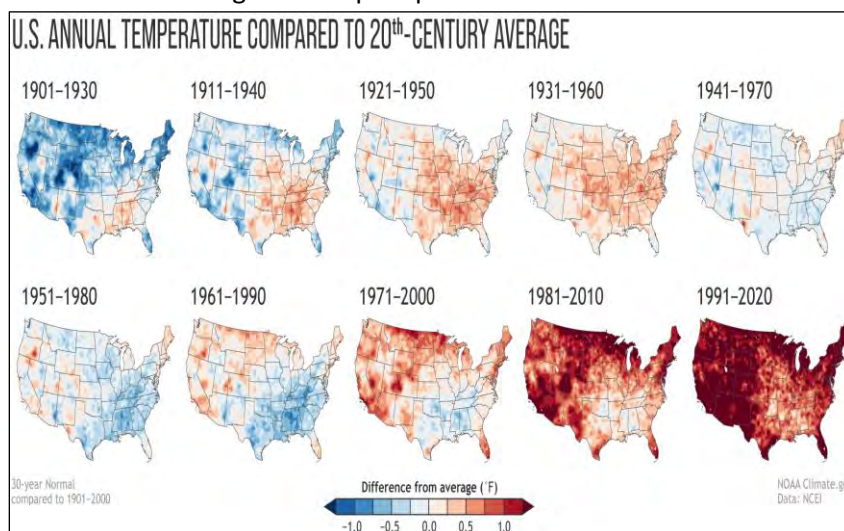


Figure 1 U.S. Annual Temperature compared to 20th-Century Average

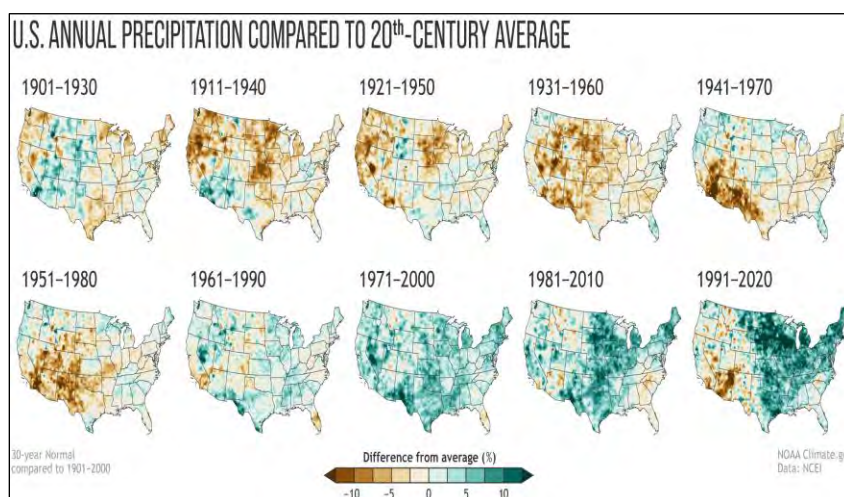


Figure 2 U.S. annual precipitation compared to 20th-Century average.

1. INTRODUCTION

Due to high natural precipitation variability and the increasing demand for water, cloud seeding has been conducted in some parts of Utah for well over 40 years. Cloud seeding in Utah is regulated by the Utah Department of Natural Resources through the Division of Water Resources. After complying with various permit requirements, NAWC was granted a license and permit to conduct cloud seeding for the High Uintas Program watersheds.

The Duchesne County Water Conservancy District and the Uintah Water Conservancy District were joined by the Central Utah Water Conservancy District in supporting an operational seeding project beginning in the 2002-2003 winter season. The state of Utah, through the Utah Division of Water Resources and the Lower Colorado River Basin States (LBS) collectively reimburse up to 50% of the cost of this program, and provide funding for an extension period that permits the continuation of the program through the month of April.

The intended target area of this program has been only the south slope of the Uinta Mountains above about 8,000 feet. However, a recent feasibility analysis and discussions with the Division of Water Resources resulted in the addition of the north slope of the Uintas (on the Utah side of the state line) to the program beginning in the 2021-2022 season. For the 2021-2022 water year, the LBS and the UDWR also provided funds for the acquisition of a remotely operated ground seeding generator that was placed at Moon Lake. This generator releases higher concentrations of seeding solution than manual generators and permits more regular and consistent operation during the winter months when the road leading up to Moon Lake is frequently closed due to snow pack.

The High Uintas Program is tributary to the Colorado River via the Green River, and LBS funds have been used to augment the program beginning in the 2010 water year. The extension period funded by Lower Basin States has been at the beginning of the core project season for the High Uintas, during the month of November each season. The extension provides additional benefit to the primary project sponsors at no additional cost to them. As additional LBS funding benefits, additional ground-based silver iodide generators have previously been added to the program, as well as strategically-located mountain ridge ice detector systems designed to help identify storm periods producing supercooled liquid water which is the target of the cloud seeding efforts. This included an ice detector which was added to the north slope of the Uintas this season, at the Lily Lake SNOTEL site due to additional available funding, also provided by the LBS.

This report provides information about operational cloud seeding conducted over the target watersheds in the 2021-2022 winter season, including the extension period. Section 2.0 describes the seeding project design and provides maps of the seeded target areas, as well as the locations of the ground-based seeding units (generators) with which the seeding was conducted. Section 3.0 describes the meteorological and computer forecast model data used in the conduct of operations, with some examples presented. Section 4.0 summarizes the seeding operations and documents the seeding generator usage by site and storm event. Section 5.0 provides an overview of statistical evaluations of the effects of the cloud seeding program.

2. PROJECT DESIGN

2.1 Background

The general project design utilized for the High Uintas cloud seeding project is essentially the same as that which has been shown to be effective for over four decades of wintertime cloud seeding in other mountainous regions of Utah. Estimations of seeding effectiveness for long-standing operational seeding projects in Utah have consistently indicated increases in winter season precipitation and snow water content during the periods in which cloud seeding was conducted. The increases for most ground-based programs have averaged approximately 5-10% more than what would have been expected in the absence of seeding, as predicted by historical target/control linear regression analyses.

The target area for the High Uintas project is adjacent to the target area for the Upper Weber Basin (Western Uintas) Project as seen in (Figure 2.1), which has also been conducted for a number of recent winter seasons. As discussed earlier the program was expanded for the 2021-2022 water year to incorporate the Northern Slope of the Uinta Mountain Range.

The target area was designed to include elevations of 8000 feet MSL or greater on the south slope of the Uinta Mountains containing river drainages that provide water to either of the sponsoring counties, plus areas providing runoff into Strawberry and Currant Creek Reservoirs. To overcome the impacts of the thermal inversions that is common in the valley south of the Uinta Mountain Range during the winter, NAWC recommends placing ground generators at or above about 7,000 feet in elevation. Due to the prevalence of national forests and Native American Reservations at upper elevations the placement of ground generators above the inversion is no simple task. This was the primary motivation for the LBS to sponsor a remote generator in the Uinta Range.

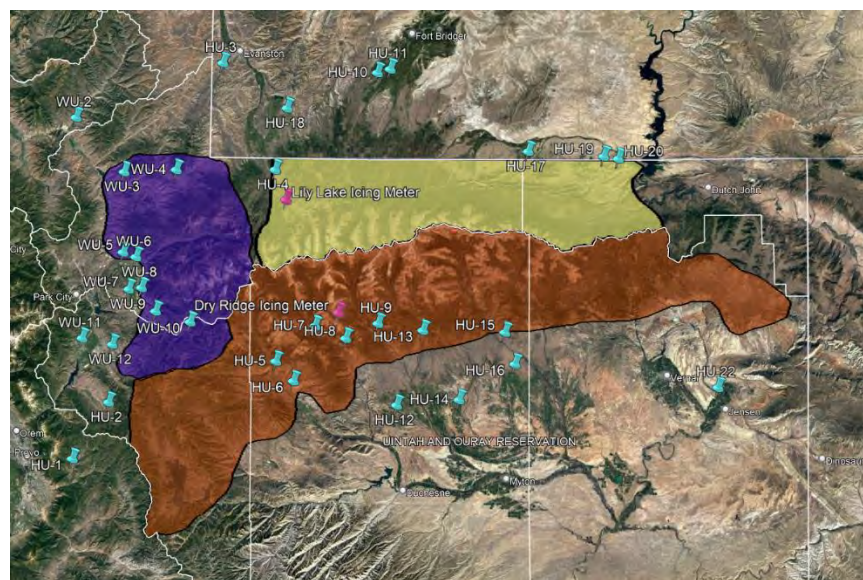


Figure 2.1 Western Uinta Program (purple), High Uinta Program (orange), Northern Slope Extension (yellow)

Regarding the second factor, project duration, Table 2-1 shows average monthly precipitation amounts at three high elevation NRCS SNOTEL sites located within the target area. The month of April is obviously a

very productive period based on climatology. Such information was used in specifying the cloud seeding project core operational period.

Table 2-1
Average Monthly Precipitation in the Target Area (inches)

Site	Elev.	Oct	Nov	Dec	Jan	Feb	Mar	Apr	May
Chepeta	10,300	2.6	2.2	2.2	2.3	2.2	2.6	3.6	2.9
Five Pts.	11,000	2.9	2.3	2.8	2.5	2.2	2.8	2.7	2.9
Trout Cr.	9,400	1.7	1.8	1.7	1.8	2.0	2.5	2.6	2.3

Consideration of the third issue (wind direction) dictates that a significant number of generators should be placed at south flank locations, since a number of the more productive storms have steering level winds from the southeast through west-southwest directions. Another maximum in potentially seedable storms occurs during westerly to west-northwesterly winds, which supports frequent usage of sites on the western side of the Uinta Range. Some seedable situations involve winds with a more significant northerly component (i.e. from northwesterly to northeasterly), and this supports the location of seeding sites on the northern side of the Uinta Range. Operational experience with this program has shown that storms with northerly-component winds may be good seeding candidates, with the enhanced snowfall on the northern slope of the Uintas that frequently carries over to the upper portion of the southern slope (within the target area) as well.

2.2 Seedability Criteria

NAWC has historically followed a selective seeding approach. This has proven to be the most efficient and cost-effective method, and provides the most beneficial results. Selective seeding, or seeding only storms or storm periods in which precipitation has a reasonable chance of being enhanced, is based on several criteria which determine the seedability of the winter storms. These criteria deal with the structure of the airmass (temperature, thermodynamic stability, wind flow and moisture content), both in and below the precipitating clouds. The following list provides a summary of the generalized criteria that NAWC uses in the conduct of its wintertime projects in the intermountain west. These criteria are based upon the results obtained in a number of relevant research-oriented weather modification programs.

NAWC Winter Cloud Seeding Criteria

- Cloud bases are near to (ideally) below the mountain barrier crest.
- Low-level wind directions and speeds would favor the movement of the silver iodide particles from their release points into the intended target area.
- No low-level atmospheric inversions or stable layers that would restrict the upward vertical transport of the silver iodide particles from the surface to at least the -5°C (23°F) level or colder.
- Temperature at mountain barrier crest height expected to be -5°C (23°F) or colder.
- Temperature at the 700mb level (approximately 10,000 feet) expected to be warmer than -15°C (5°F).

2.3 Equipment and Project Setup

During the off-season, the ground-based generators are routinely removed from the field for maintenance and testing. NAWC began re-installing the generators in October 2021. The generators were placed at the locations shown in Figure 2.1.

The cloud seeding equipment at each manually operated site consists of a cloud nuclei generator (CNG) unit and a propane gas supply. The seeding solution contains two percent (by weight) silver iodide (AgI), the active seeding agent, complexed with very small portions of sodium iodide and para-dichlorobenzene in solution with acetone. A paper published by Dr. William Finnegan, a well-respected cloud seeding formulation expert of the Desert Research Institute (Finnegan, 1999), indicates that this formulation is superior to others that produce pure silver iodide particles. The modified particles produced by combustion of the revised formulation act as ice nuclei much more quickly, and there are somewhat larger numbers of effective nuclei at warmer temperatures (i.e., about -5°C to -10°C). Figure 2.2 is a photograph of a manually operated, ground-based cloud nuclei generator such as those used for the High Uintas Program. Trained local operators are available to activate each seeding site upon request from a NAWC meteorologist. A cloud nuclei generator is activated by igniting a propane flame in the burn chamber, and then adjusting the flow of seeding solution through a flow rate meter. The propane gas also pressurizes the solution tank, which allows the solution to be sprayed into the burn chamber at a regulated rate, where microscopic-sized silver iodide (AgI) crystals are formed. The crystals, which closely resemble natural ice crystals in structure, are released at a rate of 8 grams per hour when the 2% (AgI by weight) solution is used. These crystals become active as artificial ice forming nuclei at in-cloud temperatures between -5°C and -10°C (23°F to 14°F).

It is necessary that the AgI crystals become active in supercooled clouds at relatively low altitudes upwind of (or over) the mountain crest. This allows the available supercooled liquid water to be effectively converted to ice crystals which can grow to snowflake size and precipitate onto the mountain barrier. If the AgI crystals take too long to become active, or if the temperature upwind of the crest is too warm, the plume will pass from the generator through the precipitation formation zone and over the mountain crest without producing any additional snowfall in the intended target area.



Figure 2.2 NAWC manually operated cloud nuclei generator (CNG)

Manually operated CNGs are maintained at 20 locations specific to the High Uintas program, with in addition to the newer remotely operated seeding equipment at Moon Lake on the southern flank of the target area. There are now eight sites on the northern side of the target area. Two other sites are used primarily to target the Strawberry Divide area (sites HU1 and HU2), with many of the nearby Western Uintas sites utilized to target this area as well. The network of sites is designed to be effective in generating plumes of seeding material which will pass over the target area in a variety of wind flow situations. A good number of sites primarily designated for use in the Western Uintas Program (W prefix in Figure 2-1) are also utilized for seeding the High Uintas target area when conditions are favorable for this. Pertinent site information is listed in Table 2-2, corresponding to the site numbers shown in Figure 2.1.

Table 2-2
Cloud Seeding Generator Sites

Site ID	Site Name	Elevation (Feet)	Latitude (N)	Longitude (W)
HU1	Hobble Creek	5870	40°12.22'	111°30.14'
HU2	Wallsburg	6175	40°20.95'	111°23.00'
HU3	Evanston	7000	41°12.70'	111°01.13'
HU4	Bear River East	8223	40°56.54'	110°50.17'
HU5	Hanna Pump House	7019	40°27.60'	110°49.56'
HU6	Hanna	6781	40°24.64'	110°46.03'
HU7	Rock Creek Ranch	7988	40°33.02'	110°41.78'
HU8	Robbins Ranch	7404	40°31.18'	110°35.64'
HU9	Moon Lake (remote)	8100	40°33.25'	110°29.20'
HU10	Black's Fork	7509	41°11.39'	110°29.87'
HU11	Robertson	7322	41°11.97'	110°27.31'
HU12	Talmage	6945	40°21.53'	110°27.28'
HU13	Yellowstone Canyon	7660	40°32.50'	110°20.30'
HU14	Bluebell	5840	40°26.85'	110°03.72'
HU15	Uinta Power Plant	6932	40°32.27'	110°03.98'
HU16	Neola	6330	40°27.48'	110°02.93'
HU17	Birch Creek	7634	40°58.64'	109°59.48'
HU18	Gilmore Ranch	7550	41°05.91'	110°47.96'
HU19	Manila	6500	40°58.91'	109°44.36'
HU20	Manila East	6230	40°58.57'	109°41.38'
HU22	Jensen	4896	40°23.92'	109°21.49'
W4	Pineview	6407	40°56.39'	111°10.18'
W6	Oakley	6472	40°43.07'	111°18.00'
W7	Kamas	6489	40°38.43'	111°16.77'
W8	Kamas West	6472	40°38.16'	111°19.33'
W9	Woodland	6706	40°34.89'	111°13.81'
W10	Woodland East	7305	40°33.35'	111°06.80'
W11	Midway	5570	40°30.59'	111°28.64'
W12	Heber City	5810	40°29.73'	111°22.52'

2.4 Project Instrumentation

Some specialized instrumentation has been added over the past number of years to enhance cloud seeding guidance during operations within the High Uinta Program area. This includes icing rate meters and, during one previous (2019-2020) season, a radiometer was located in the Uinta Basin. Both instrument systems were supported by funding from the Lower Basin States. For the 2021-2022 season, a second icing rate meter was used was installed at Lily Lake, on the northwestern side of the Uinta Range at an elevation of 9100'. This provides a comparison to a similar site located at Dry Ridge on the southern slope, although the Dry Ridge site is quite a bit higher in elevation at over 11,500'. Both of these sites provide real-time observations of supercooled liquid water (SLW). Other observations include precipitation, temperature, wind direction and wind speed. Because SLW is the target of cloud seeding, such a sensor is of benefit both in terms of real-time operational decisions and for later analysis of the frequency of SLW occurrence in relation to winter storm periods. This sensor is similar to sensors which have been installed in other seeding target areas in Utah. Analysis reports on the Utah ice detector data are available on the NAWC website at <http://www.nawcinc.com/publications.html>. Analyses of the data from these sites have provided valuable insight into the occurrence of SLW during winter storms. Figures 2.3 and 2.4 provide photographs of the installation. The funding for the equipment, installation and maintenance of this site was provided by three Lower Colorado River Basin States and administered by the Utah Department of Water Resources Division.



Figure 2.3 Icing Rate Meter Installation at the Dry Ridge Site



Figure 2.4 Dry Ridge Sensor Suite



Figure 2.5 Icing Meter at Lily Lake SNOTEL Site

2.5 Suspension Criteria

NAWC always conducts its projects within guidelines adopted to ensure public safety. Accordingly, NAWC has a standing policy and project-specific procedures for the suspension of cloud seeding operations in certain situations. Those criteria are shown in Appendix A, and have recently been updated in coordination with the Utah Division of Water Resources. The criteria are an integral part of the seeding program. There were no suspensions of seeding operations during the 2021-22 season.

3. WEATHER DATA AND MODELS USED IN SEEDING OPERATIONS

NAWC maintains a fully equipped project operations center at its Sandy, Utah headquarters. Meteorological information is acquired online from a wide variety of freely available sources and subscriber services. This information includes weather forecast model data, surface observations, rawinsonde (weather balloon) upper-air observations, satellite images, NEXRAD radar information, and weather cameras. This information helps NAWC's meteorologists to determine when conditions are appropriate for cloud seeding. NAWC's meteorologists are able to access all meteorological products from their homes, allowing continued monitoring and conduct of seeding operations outside of regular business hours.

Figures 3.1 – 3.3 show examples of some of the available weather information that was used in this decision-making process. Figure 3.4 provides predictions of ground-based seeding plume dispersion for a discrete storm period in the High Uintas using the National Oceanic and Atmospheric Administration's HYSPLIT model. This model helps to estimate the horizontal and vertical spread of a plume from potential ground-based seeding sites in real-time, based on wind fields contained in the weather forecast models.

Global and regional forecast models are a cornerstone of modern weather forecasting, and an important tool for operational meteorologists. These models forecast a variety of parameters at different levels of the atmosphere, including winds, temperatures, moisture, and surface parameters such as accumulated precipitation. An example of a display from the Global Forecast Systems (GFS) model is shown in Figures 3.5.

Figure 3.6 is an example of a plot from the radiometer installed on the north slope of the Uintas this season, near the Lily Lake icing meter site. The radiometer is a passive instrument that can measure several meteorological variables, displayed in cross-sectional plots as shown in this figure.

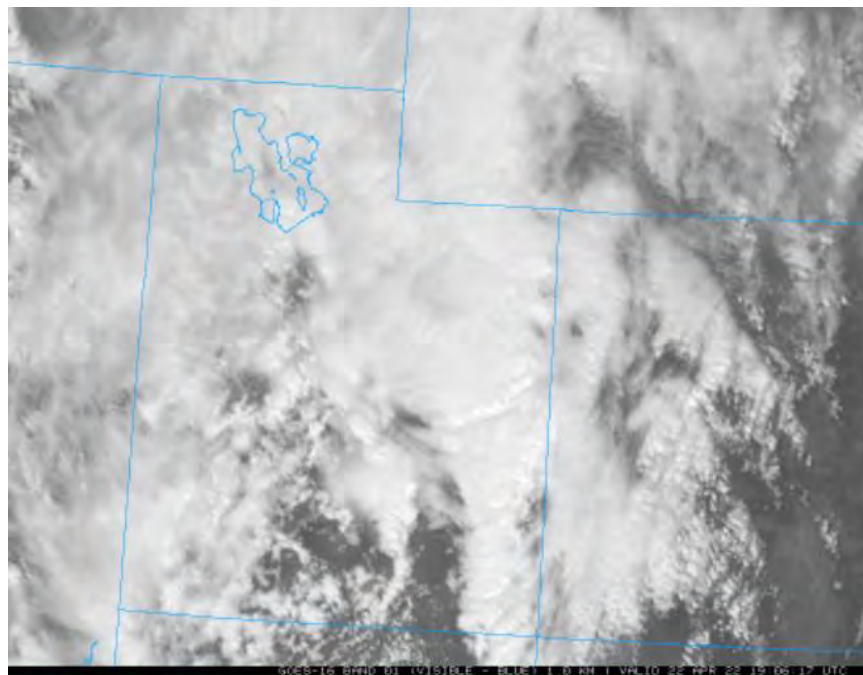


Figure 3.1 Visible spectrum satellite image on March 22, 2022 during a seeded event.

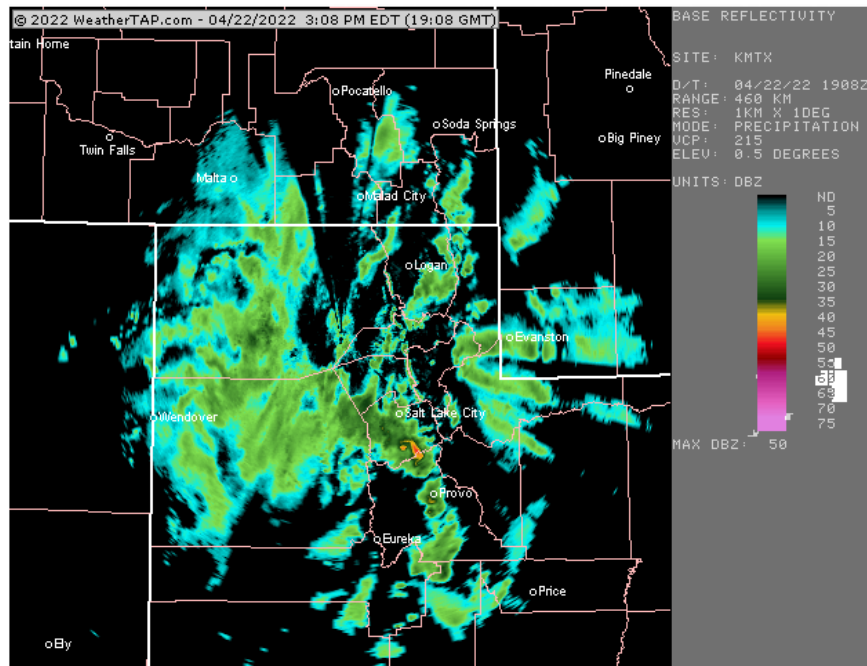


Figure 3.2 Weather radar image over northern Utah on April 22, corresponding to the satellite image in Figure 3.1

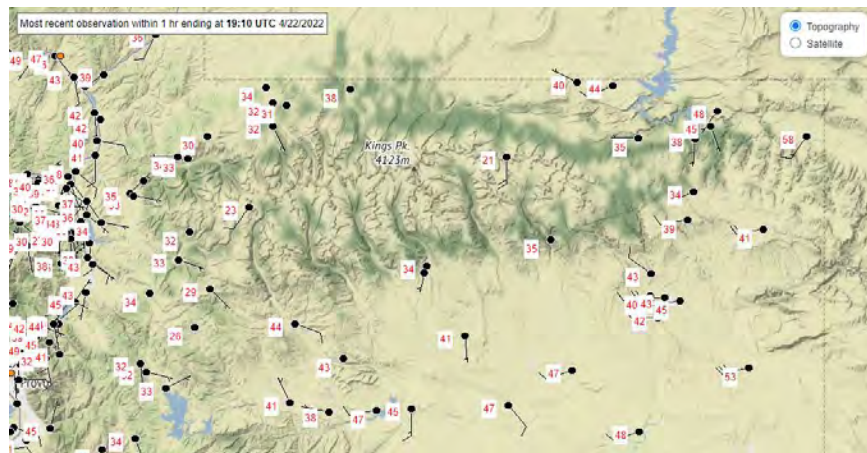


Figure 3.3 Surface map during the April 22 storm event, showing winds and temperatures at a variety of automated sites in an around the Uinta Range.

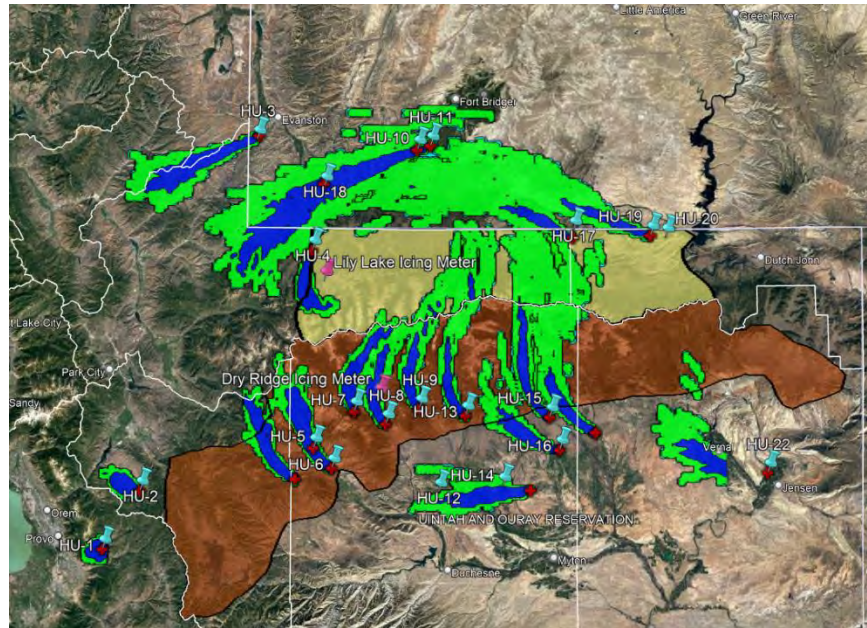
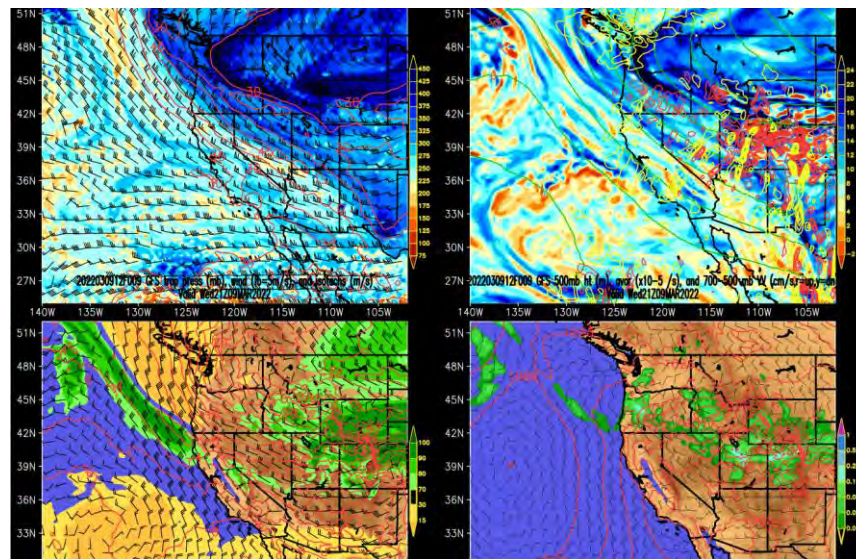


Figure 3.4 HYSPLIT plume dispersion forecast for potential seeding locations during a storm on the evening of March 5, 2022. This is a tool that can be used to help select appropriate sites for a given situation.



Figures 3.5 GFS model 4-panel data display during a storm event on March 9, 2022. The lower left panel shows winds, moisture, and temperature at the 700-mb level which are especially useful for seeding operations.

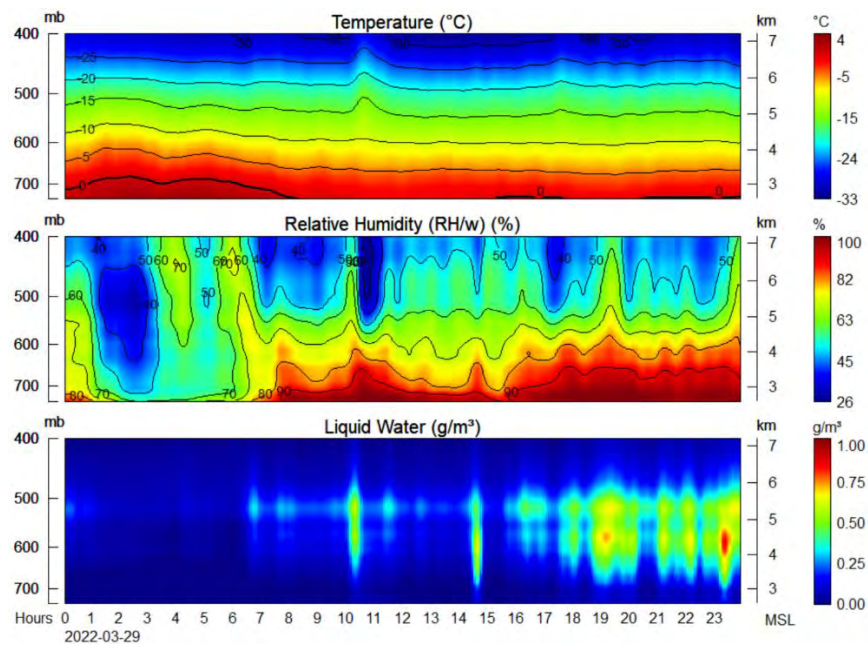


Figure 3.6 Radiometer cross sections (temperature, humidity, and liquid water) measured in a storm event on March 29, 2022

4. OPERATIONS

The core 2021-2022 cloud seeding program for the High Uintas contractually extended from December 1, 2021 through April 30, 2022, with an extension period from November 1-30, 2021 funded by the Lower Basin States. During the entire operational season of November 1 – April 30, seeding operations took place over 23 storm periods, with two of these occurring during the extension period in November. Altogether, there were two seeded storms in November, five in December, two in January, two in February, seven in March, and five in April. A cumulative 1,617.5 hours of ground seeding generator operations were conducted during the regular season, and an additional 115 hours during the extension period, for a total of 1732.5 hours. Figure 4.1 is a graph of operations this season for the core High Uintas program, compared to a linear usage of the total budgeted hours. Table 4-1 shows the seeding dates and ground generator usage for the storm events, and Appendix B shows detailed site usage data.

Precipitation/snowfall was generally near to below average this season. As of April 1, 2022, SNOTEL observations for the Natural Resource Conservation Service showed snow water content averaging about 88% of normal (median) for the Duchesne Basin and about 91% of normal for sites in the portions of the Uinta Range that compose the Green River Basin. Water year precipitation percentages were 107% of the median for the Duchesne Basin and around 105% of normal for sites in the Green River Basin. These above average (median) values for water year precipitation were due in part to a very wet month of October, prior to any significant snowpack accumulation. By the end of the project (May 1), median snowpack percentages had decreased to 66% for the Duchesne Basin and 52% for the Green River Basin. Water year to date percentages (of the mean) on May 1 were 101% for the Duchesne Basin and 97% for the Green River Basin. Figures 4.2 to 4.4 show snow water content and water year precipitation accumulations, and normals, for October 1 through May 1 for target area SNOTEL sites.

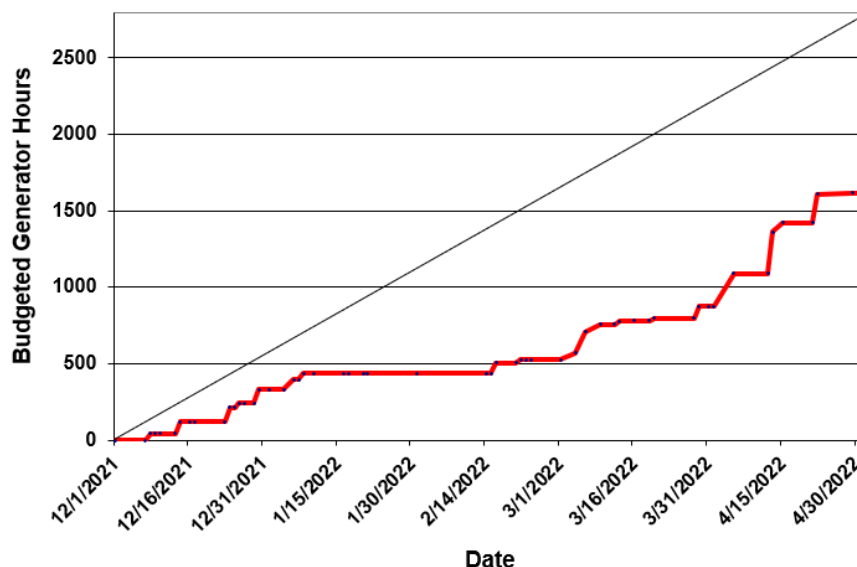


Figure 4.1 Seeding operations during the 2021-2022 season for the core program (red). Diagonal black line shows a linear usage of total budgeted hours, as a reference.

Table 4-1
Storm Dates and Number of Generators used in the High Uintas Program

Storm Number	Date	Number of Generators	Operational Hours
1*	November 9-10	8	101.25
2*	November 19-20	2	13.75
3	December 8-9	6	44.5
4	December 14-15	9 + remote	73.75 + 6 remote
5	December 24-25	5	96.75
6	December 26	5 + remote	25 + 10.75 remote
7	December 30-31	7	91.5
8	January 5-6	3	66
9	January 7-8	3	39
10	February 16	7	67
11	February 21	4	20
12	March 4	9 + remote	47.5 + 2 remote
13	March 5-6	12 + remote	139 + 3 remote
14	March 8-9	4	48.75
15	March 13	3	19.25
16	March 16	1	4.75
17	March 20	2	12.75
18	March 29	9	80.75
19	April 4-5	14	212
20	April 11-12	14	275
21	April 14-15	4	61
22	April 22-23	15	186.25
23	April 29	2	7
Core Program Total			1617.5
Extension Total			115

* Seeding for Lower Basin-Funded Extension

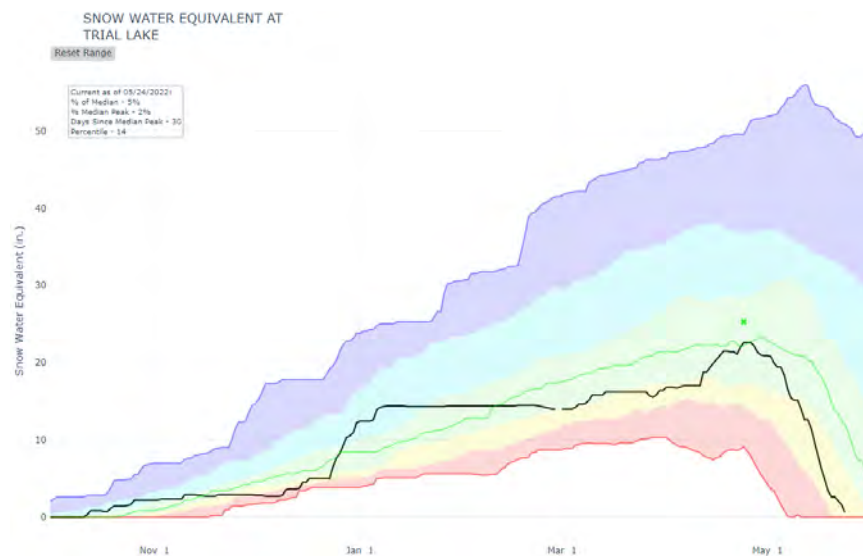


Figure 4.2 NRCS SNOTEL snow water content plot for October 2021 through May 2022 for the Trial Lake SNOTEL, UT in the western Uintas. Black line is the 2021-22 season data. Green represents the median, and purple and red are the historical maximum and minimum values respectively.

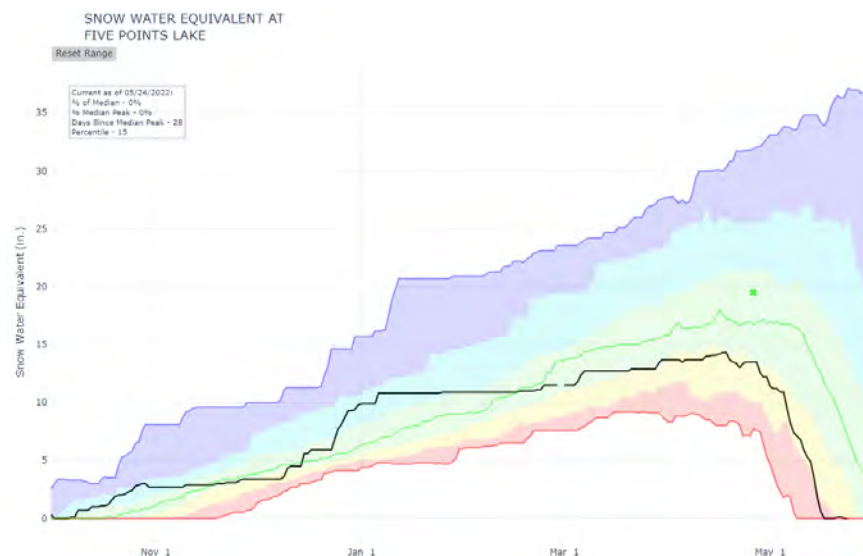


Figure 4.3 NRCS SNOTEL snow water content plot for October 2021 through May 2022 for the Five Points Lake SNOTEL, UT in the central portion of the Uintas. Black line is the 2021-22 season data. Green represents the median, and purple and red are the historical maximum and minimum values respectively.

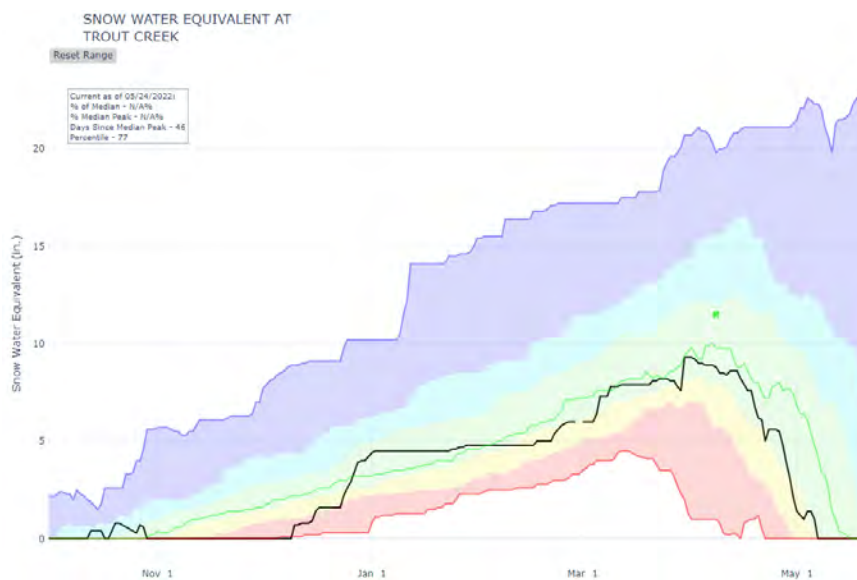


Figure 4.4 NRCS SNOTEL snow water content plot for October 2021 through May 2022 for the Trout Creek SNOTEL, UT on the eastern side of the Uintas. Black line is the 2021-22 season data. Green represents the median, and purple and red are the historical maximum and minimum values respectively.

4.1 Operational Procedures

In operational practice, the project meteorologist, with the aid of continually updated online weather information, monitored each approaching storm. If the storm parameters met the seedability criteria presented in Table 2-1, and if no seeding curtailments or suspensions were in effect, an appropriate array of seeding generators was ignited and then adjusted as evolving conditions required. Seeding continued as long as conditions were favorable and precipitating clouds remained over the target area. The operation of the seeding sites is not a simple “all-or-nothing” situation. Individual seeding sites are selected and run based on their location, and targeting considerations based on storm attributes.

4.2 Operational Summary

A brief synopsis of the weather during the operational seeding period is provided below. All times reported are local, either in MST or MDT. When wind direction information is given it is the direction from which the wind is blowing. For example, a northwest wind is blowing from the northwest towards the southeast. The temperature at the 700 mb level (~9,500 feet above sea level during the winter) is commonly referenced, since temperature is an important factor when determining the seeding potential of an event. Data from the ice detector site at Dry Ridge (elevation 11,540 feet) can also be an important indicator of the presence of supercooled water in the target area, and thus seeding potential.

November 2021

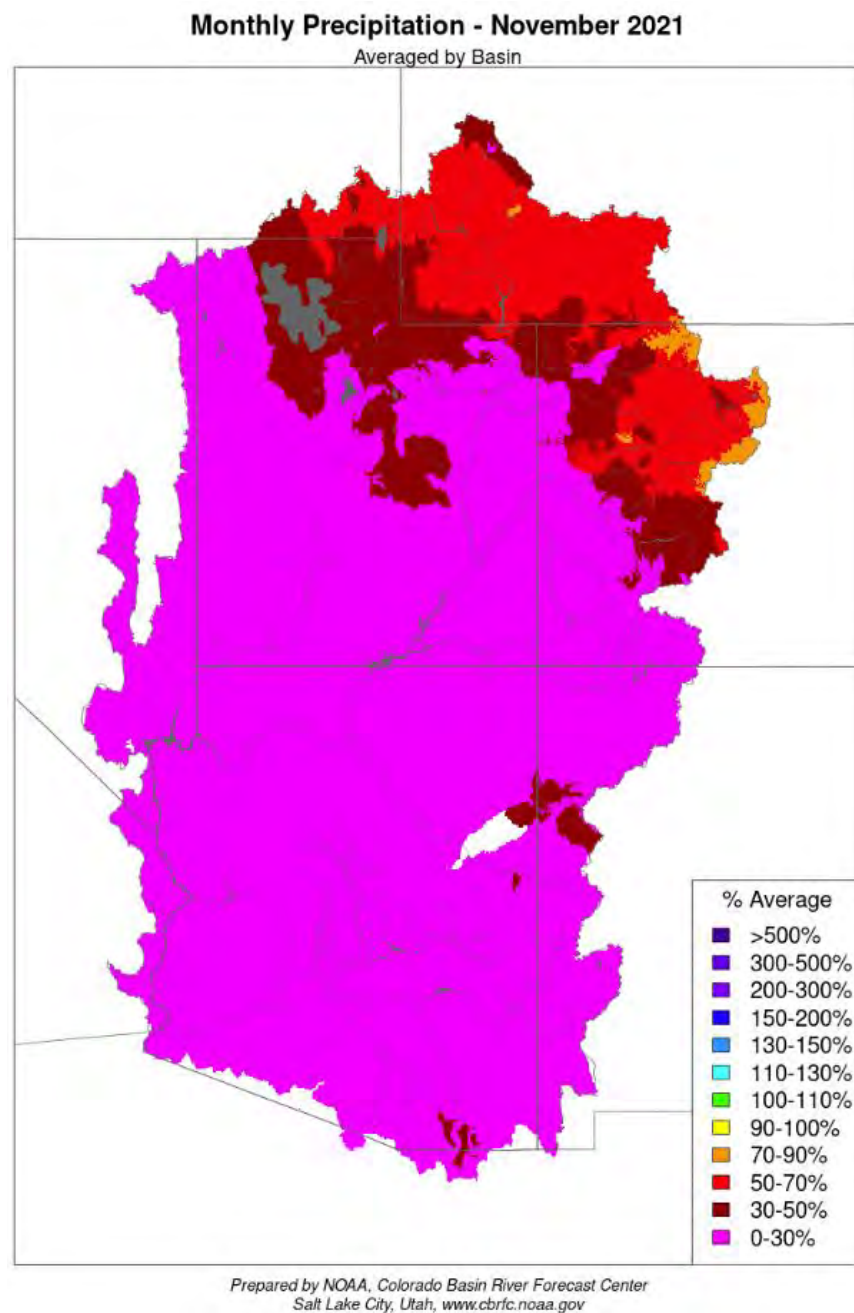


Figure 4.5 November 2021 precipitation, percent of normal

December 2021

Following the alternating pattern of the previous several months, December was a very wet and stormy month with the Uinta Range receiving generally over twice the normal monthly precipitation and snowfall. Most of this activity was later in the month with a series of significant storms during the final week of December. In total, there were five seeding storm periods in December.

A cold front on the night of December 8-9 brought a period of seeding opportunity using sites on the west side of the Uinta Range. The 700 mb temperature cooled to around -5° C with westerly to southwesterly winds. There were indications of orographic (terrain induced) precipitation in radar imagery and also indications of at least patchy areas of liquid water development, necessary for effective seeding. Dry Ridge recorded one icing cycle on the evening of the 8th and a couple others on the 9th. Seeding operations continued until late morning on December 9 when it became apparent that clouds were all of the higher, icy variety at that point, and seeding operations ended. Seeding operations utilized mostly sites in and near the target area on the west/southwest side in this event. Precipitation totals were generally in the 0.5 – 1.0” range (water content).

A cold frontal passage on the night of December 14-15 produced a good seeding opportunity. Some low level moisture streaming up from the south ahead of the front produced clouds fairly rich in liquid water by the evening of the 14th, although temperatures were fairly mild ahead of the front (around -2 C at 700 mb). A good amount of icing was observed at the Dry Ridge site during the afternoon and evening, with a total of 39 icing cycles there in the early stages of this storm period. The Lily Lake site on the northern slope did not record any icing. Seeding began later in the evening and continued overnight as the frontal band moved through with a period of moderate to briefly heavy snowfall. The 700 mb temperature cooled to near -12° C behind the front. Seeding ended early on the 15th with drying conditions. This storm event brought about 0.5 – 1.0” of water content to the area, with seeding again conducted mainly from sites on the southwestern side of the Uintas.

A series of storms which began on December 23 brought snowfall from mainly a high cloud deck initially, not favorable for seeding operations. Although there were a few icing cycles at Dry Ridge in the early morning hours of the 23rd prior to the onset of this event, there was no other icing at either site that day. By the 24th the situation changed with convective showers and even some thundershowers developing during the late afternoon and evening hours in southwesterly flow. Seeding was conducted from sites on the south and southwestern side of the Uintas based on wind direction. The seeding continued into the morning to midday of the 25th before ending, as a brief period of drying occurred. Precipitation total of 0.2 to 0.8” were common during the seeded storm period. There was no icing observed at the sites during this storm period.

A fast-moving cold front affected the area on December 26, keeping widespread snowfall across the area. Temperatures were much colder than previously, falling from about -8 to -14° C and liquid water appeared much more limited. However, seeding was conducted from a few sites in the target area during the daytime hours. Precipitation totals for this period as a whole ranged from about a half inch to just over an inch. Once again, no icing was observed at Dry Ridge or Lily Lake.

Frequent snowfall continued through the last few days of December, although much of it was associated with cold temperatures and a general lack of liquid water with ice processes being very dominant. There was a system at the end of the month that was determined to be favorable for effecting seeding, beginning on the afternoon of December 30. The 700 mb temperature warmed to around -8°C with southwesterly flow in the warm sector of this system, and there were various indication of liquid water clouds favorable to seeding. Seeding was conducted from sites on the southern and southwestern side of the Uintas, and continued overnight. Winds became northwesterly with much colder arctic air moving into the area on December 31, with seeding ending during the morning hours. There was no icing observed at the Dry Ridge or Lily Lake detector sites during these late December storm periods. The seeded storm of December 30-31 period produced amounts ranging from about 0.2 to 0.9" of water content. Figure 4.6 shows December 2021 precipitation as a percentage of the median across the region.

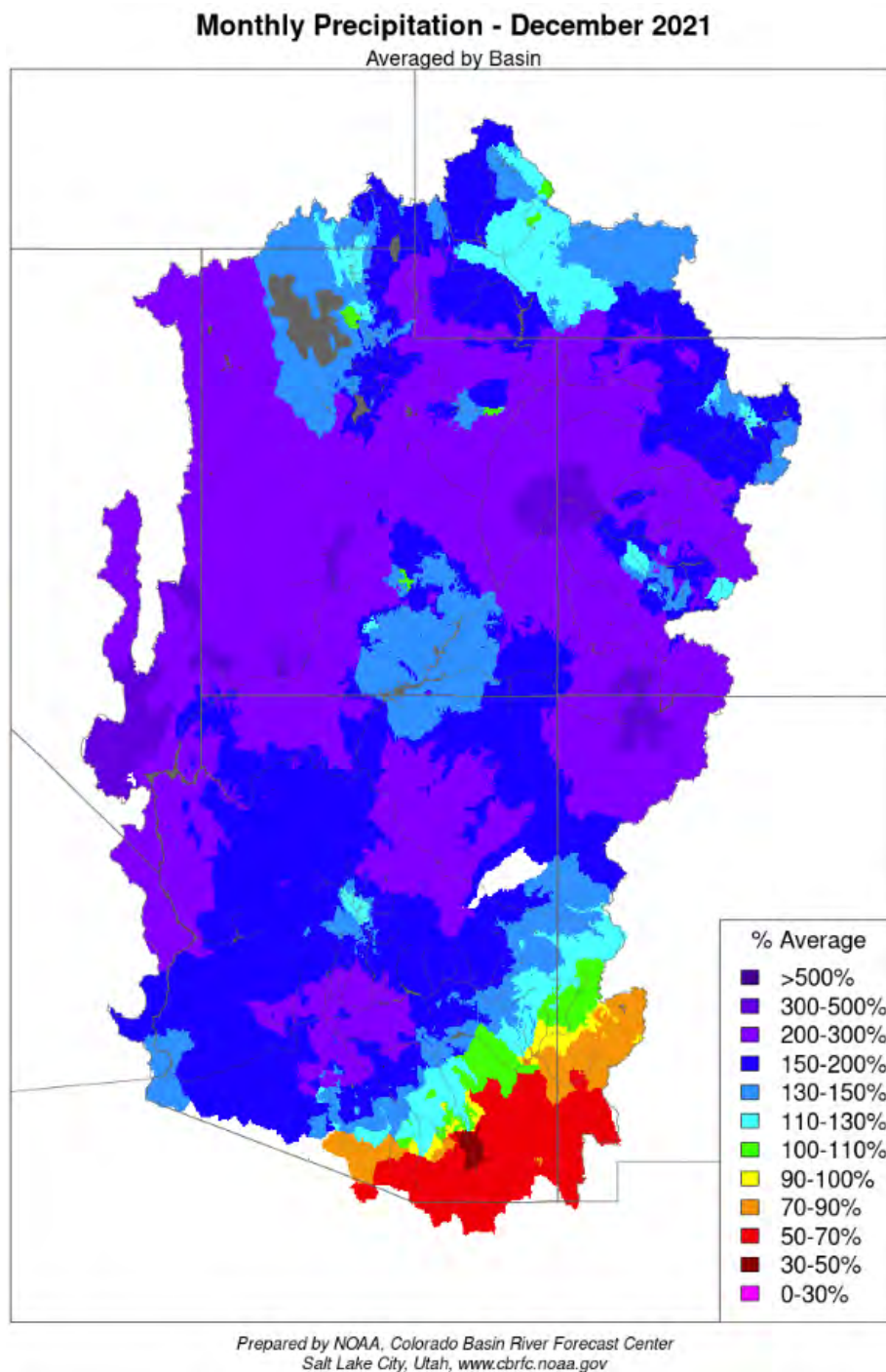


Figure 4.6 **December 2021 precipitation, percent of normal**

January 2022

Except for an extended storm period affecting northern Utah early in the month and providing a couple of seeding opportunities, January was generally quite dry with only weak storms reaching the area. Also many of these storms were cold/icy or otherwise unfavorable for seeding operations. Precipitation for the month as a whole was well below average, and what did occur strongly favored the northern side of the Uintas.

A significant moisture plume from the Pacific affected the area during the January 5-7 period, in a generally west-northwesterly wind pattern. Seeding was conducted to target mainly the northern and western side of the Uintas from western sites on the night/morning of January 5-6, with 700 mb temperature generally near -5°C although warming to around -3°C during the day on the 6th. There were indications of at least sufficient liquid water for seeding operations during this time, and one icing cycle was observed at Dry Ridge overnight (with more on the afternoon of the 6th after temperatures had warmed). The Lily Lake site, as in most events, did not record icing. Seeding ended with some brief drying on much of January 6, with precipitation totals ranging from generally 0.6 – 1.3" at SNOTEL sites during this first storm period.

Another period of seeding was conducted on the night of January 7-8 as a cold front on the tail end of the Pacific moisture plume crossed the area. Moisture was a bit lacking in this portion of the system, but seeding was conducted from a few sites on the western end of the Uinta Range as temperatures cooled to well below -5°C at 700 mb by early on the 8th. There was one icing cycle at Dry Ridge on the night of January 7-8. Seeding ended early with partial clearing, and precipitation totals were in the 0.1 – 0.4" range from this system.

For the remainder of January, a combination of widespread temperature inversions, some arctic air masses and weak systems with limited moisture (and virtually no liquid water apparent) precluded any additional seeding operations. Figure 4.7 shows the January precipitation as a percentage of the monthly median, with total generally over 50% of the monthly average on the north side and well under 50% of the normal amounts on the southern side of the Uintas.

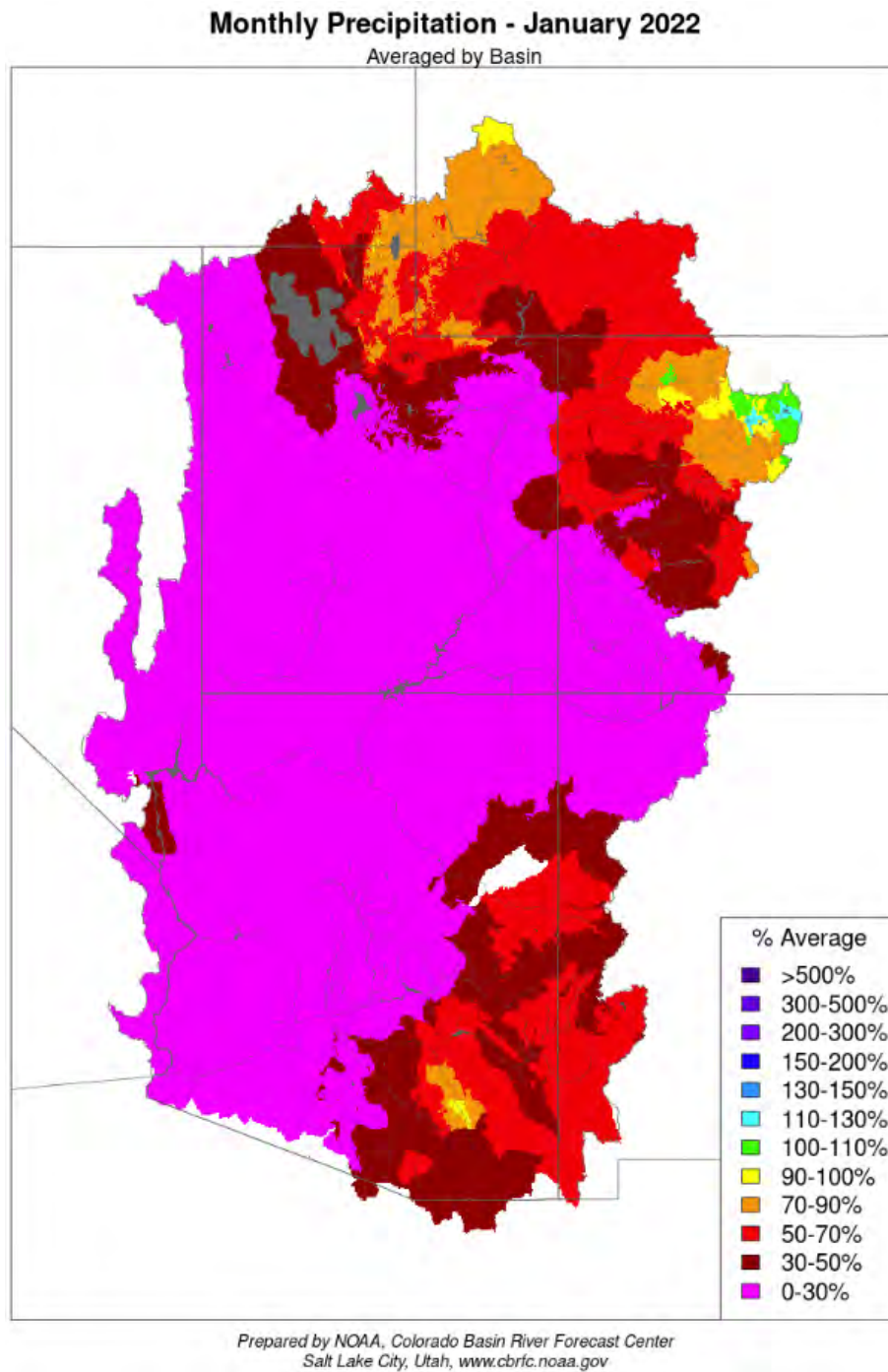


Figure 4.7 **January 2022 precipitation, percent of normal**

February 2022

The month of February continued the same general pattern as January, being very dry with extensive valley temperature inversions over the region at times. There were a few systems that moved from north to south across the area, generally lacking in liquid water clouds and favoring the northeastern portions of the Uintas. There were only two seeding opportunities in February.

In addition to the data from the ice detector sites at Dry Ridge and Lily Lake, data from a radiometer near Lily Lake is available for storm events beginning in February this season. This data is valuable for helping to confirm when (and how much) cloud liquid water was available for seeding.

A trough moving from north to south across the area on February 16 brought some snow showers during the daytime hours as temperatures cooled to around -12°C at 700 mb. Seeding was conducted during the day from north-side sites as winds were generally from the north to northwest, and snow showers appeared to have at least limited amounts of liquid water as confirmed by radiometer data. Dry Ridge had one icing cycle prior to this period, during the night of February 15-16. Precipitation totals were generally around a quarter to half inch of water in the target area.

Some limited seeding was conducted on February 21 around the east and northeast side of the Uintas, with a frontal zone and band of snowfall over the area. Wind patterns were complex as arctic air began to arrive from the northeast, and much of the snowfall was from a higher cloud deck with limited liquid water. Precipitation totals were on the light side with generally 0.1 – 0.3" of water content. The Lily Lake ice detector site indicated 4 cycles in one period around midday on February 21, although based on observed patterns it is not currently known if this was indicative of real icing activity at the site. The radiometer near Lily Lake did show at least modest amounts of cloud liquid water, around 0.10 g/m^3 .

Figure 4.8 shows February precipitation as a percentage of average, with some portions of the Uintas receiving up to about 50% or a bit more on the eastern/northeastern side. Western portions of the Uintas received very little precipitation during February.

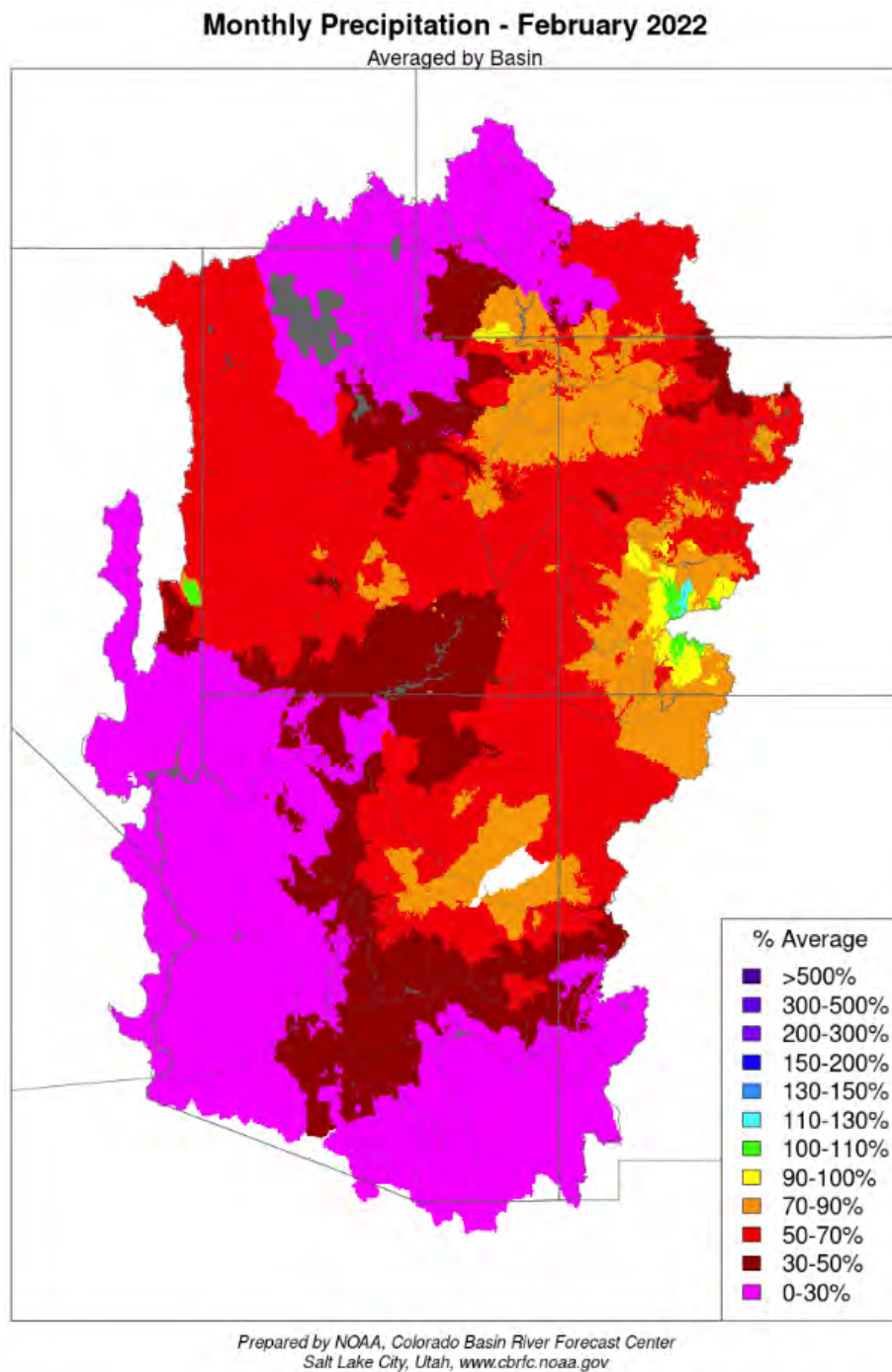


Figure 4.8 February 2022 precipitation, percent of normal

March 2022

March was a somewhat more active month with fairly frequent storm events. However, many of these were somewhat lacking in moisture and a good number only affected a portion of the Uinta Range. Precipitation totals in March were once again quite variable and strongly favored the eastern portions of the Uinta Range. Due to the fairly regular frequency of storm events in March, there were a total of seven seeded storm periods.

A disorganized trough on March 4 resulted in some convective shower activity beginning in southerly flow after midday. This activity lasted through the afternoon and had generally ended in the early evening before sunset, and, although temperatures were on the mild side (0 to -2° C at 700 mb) it appeared favorable for some seeding from south-side sites including those in the Uinta Basin. Most areas of the Uintas received a few inches of snowfall above 8,000 feet with about a quarter to half inch of water content, favoring eastern portions. There were a couple of icing cycles at Dry Ridge (one mid-morning and the other in the early evening), confirming the presence of supercooled clouds in this event. One afternoon convective shower contained LW values up to near 0.5 g/m³ for a short time period.

A fairly strong closed low had developed over the southern Great Basin by March 5, with increasing moisture in a southerly flow pattern across Utah. The 700 mb temperature was around -6° C with convective showers developing in the afternoon and evening, including a few thundershowers. Seeding was conducted from south-side sites during the day on the 5th; however, winds shifted to northerly overnight as the closed low moved into the Four Corners region, and seeding was subsequently conducted from the evening of March 5 through the morning of March 6 from north-side sites as temperatures cooled to near or below -10° C at 700 mb. The Lily Lake icing meter site recorded one icing cycle near the end of the event on March 6 when winds were in a northerly pattern, although it appeared that the southerly wind portion of this event with convective activity had by far the larger amounts of liquid water clouds. This was supported by intermittent icing observed at Dry Ridge from about midday through the early evening hours of March 5, as well as the radiometer data which showed intermittent liquid water values up to around 0.3 g/m³ at times during the afternoon and evening of March 5. Much lower liquid water values were observed on March 6. Precipitation totals during this event ranged from about 0.3 – 0.8" at SNOTEL sites in the target area.

Moisture increased in west-northwesterly flow on March 8, with lower-level winds being somewhat southwesterly. Seeding was initiated from mostly western Uintas sites to target the High Uintas program in this wind pattern. Although Dry Ridge did not record icing activity, there were two cycles at the Lily Lake site on March 8, one around midday and another during the midafternoon. This corresponded to radiometer data showing values of between 0.1 – 0.2 g/m³ during that time period. The 700 mb temperature was near -11° C. Winds became somewhat more northwesterly on the 9th with a somewhat moist, unstable air mass leading to scattered convective type snow showers. The air mass continued to cool with drying later in the day, as an arctic air mass began to arrive from the north. Seeding continued through the day from sites on the western side of the target area, ending during the evening hours. Precipitation totals were mostly in the 0.2 – 0.6" range during this event.

A fast moving cold frontal passage on March 13 was followed by the development of some convective showers in the afternoon and evening, as the 700 mb temperature fell to between -5 to -10° C. A few seeding sites on the western and northwestern side of the Uintas were able to target this activity, with seeding ending in the evening hours as skies mostly cleared. Precipitation totals ranged from about 0.1 – 0.4” at target area SNOTEL sites. The Lily Lake site recorded intermittent icing during the midday to early afternoon period on March 13, while Dry Ridge did not have any icing except one cycle during the overnight (March 13-14) period after precipitation activity with this event had ended. The radiometer data showed a distinct (but brief) period of liquid water ranging up to around 0.4 g/m³ between 1400-1500 MST on the afternoon of March 13, presumably associated with convective activity over the site.

A few light snow showers occurred in some portions of the Uintas in northwesterly flow on March 16, with one site in that area targeting these showers with some seeding during the afternoon. This was a very minor event with only about 0.1” of water content in affected areas, and no icing activity was recorded.

Another minor seeding opportunity occurred on March 20 with snow showers during the afternoon and early evening following a cold frontal passage. A few sites on the west and northwestern side of the Uintas were utilized based on wind direction. The 700 mb temperature dropped to about -10° C during this period and precipitation totals of 0.1 to 0.3” were observed. No icing activity was observed.

A large trough over the region, combined with a fairly rich lower-level moisture field and daytime heating resulting in widespread shower (and isolated thundershower) activity from midday to the early evening of March 29. The 700 mb temperature was around -3 to -4° C, suitably cold over the Uinta Range with this fairly deep convective activity. Seeding was conducted from the late morning (in anticipation of shower development) through the evening hours from various sites on the western and northern sides of the Uinta Range, with winds becoming quite northerly by later in the day. It was very notable that precipitation well exceeded forecast amounts on this day, with observed totals ranging from about 0.4 to 0.8” in western portions of the Uinta Range and some even greater amounts of about 1.0 to 1.6” observed on the eastern side. Essentially all portions of the range were targeted with at least some seeding activity on this day. The March 29 event did produce some icing activity during the afternoon hours at Dry Ridge, and there was a period with icing recorded at Lily Lake for about an hour during the early evening of March 29. Radiometer data showed widespread liquid water values in the range of 0.5 – 1.0 g/m³ from midday into the early evening hours of March 29.

Figure 4.9 shows the March 2022 precipitation totals as a percent of median, with a notable east-west gradient across the Uinta Range. Western portions of the range were well below normal for monthly totals once again, while far eastern portions of the Uintas far exceeded their March average.

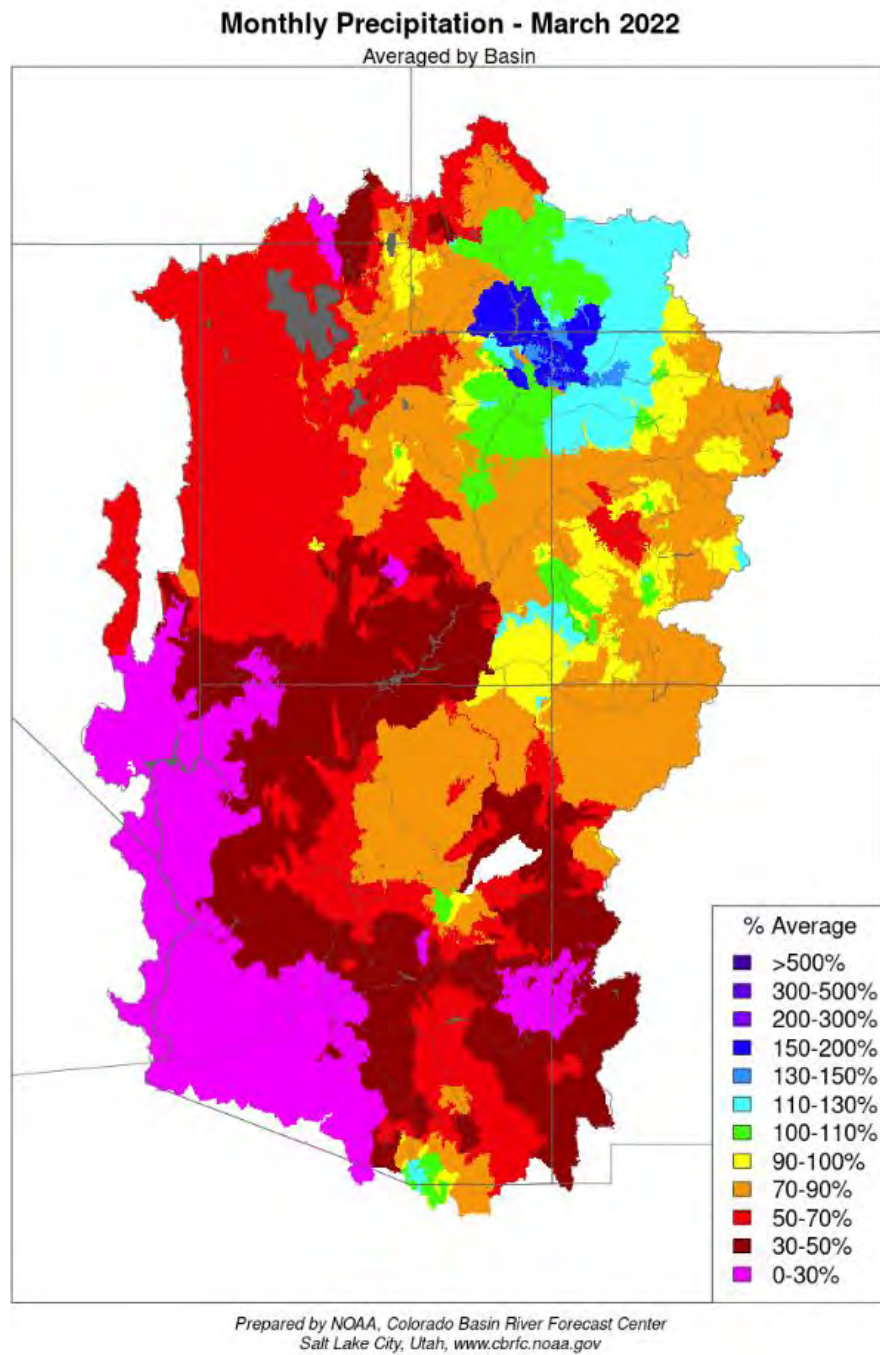


Figure 4.9 **March 2022 precipitation, percent of normal**

April 2022

April was a near normal month for some portions of northern Utah and the western end of the Uinta Range, while most surrounding areas (including central and eastern portions of the Uintas) were far below average. Storm events did occur on a fairly regular basis in April, resulting in five seeding opportunities.

A large trough on April 4-5 resulted in some snowfall across the Uintas in a westerly wind pattern. The 700 mb temperature fell from about -3 to -8° C during this time period. Seeding was conducted from nearly all sites on the western side and some higher sites in the target area from the evening of the 4th until almost midday on the 5th. Lower-level moisture was somewhat lacking initially with this system but there appeared to be significant seedable cloud development during the seeded portion, with good orographic effects. The Dry Ridge site recorded 4 icing cycles during the overnight hours, while Lily Lake recorded a period of icing between about 0400 – 0500 on the morning of the 5th. Radiometer data appeared to have some problems during this time period, and was hard to interpret. Precipitation was generally confined to the western side of the Uintas, ranging up to over a half inch of water content in some areas.

A strong frontal system on April 11-12 resulted in a seeding opportunity from nearly all sites on the western side of the Uinta Range. The 700 mb temperature fell from around 0° C on the 11th to about -15° C on the 12th. Dry Ridge recorded icing activity on the afternoon and evening of the 11th with a total of 7 icing cycles, while Lily Lake did not record icing. Radiometer data showed modest amounts of liquid water, up to around 0.1 g/m³ on the afternoon/evening of the 11th. The nearly due westerly wind pattern meant that the western end of the Uintas received by far the bulk of the activity with this system, with seeding beginning on the evening of April 11 and continuing through most of the 12th with scattered snow showers continuing in westerly flow during the daytime hours. Moisture became very limited by later in the day on the 12th and seeding ended at that point, with some lingering snow showers in a cold air mass on the 13th appearing too lacking in moisture for any effective seeding on that day. Radiometer data on April 12 again showed limited liquid water values, mostly on the order of 0.1 g/m³ in the snow showers during the afternoon hours. Precipitation totals during the April 11-12 storm period were fairly generous, up to around 1.2" of water content on the western side of the Uintas, but dropping off to less than 0.3" on the eastern side where only some brief periods of precipitation were observed.

A cold trough of low pressure remained over the northwestern U.S. for a few days, resulting in another minor seeding opportunity on the night of April 14-15. A weak frontal zone resulted in some snow showers in westerly flow with a 700 mb temperature near -5° C. Although moisture was somewhat lacking, seeding was conducted from several sites in and near the target area on the southwestern side. There was no icing observed at the sites during this event. Radiometer data confirmed some limited liquid water values during this time period, up to around 0.1 g/m³. Precipitation was confined to western portions of the Uintas, ranging up to about a half inch at some SNOTEL sites there.

A large trough moved from Nevada into Utah on April 22. The air mass over Utah was initially very dry with lack of any real liquid clouds initially, but moisture began to increase by midday with good mixing

and orographic lifting in a southerly wind pattern. The Dry Ridge site observed several icing cycles around midday on the 22nd (and a couple more later on, from the evening through overnight). This compares well with the radiometer data showing intermittent periods of significant liquid water, up to over 0.5 g/m^3 . Seeding was conducted from the southern side of the Uintas initially, during the afternoon hours, then adjusted to sites on the west and northwest sides during the evening and overnight as winds shifted to that direction. The 700 mb temperature was near 0° C ahead of the system and cooled to around -8° C on the night of April 22-23 and during the day on the 23rd. Snow showers continued in a cold and moist northwesterly wind pattern on April 23, with seeding conducted from appropriate sites through the day before ending in the early evening hours. These showers contained liquid water concentrations ranged from about $0.3 - 0.7 \text{ g/m}^3$ per radiometer data. Precipitation totals in the Uintas were fairly generous during this storm event, ranging from about 0.5 to 1.5" of water content at SNOTEL sites.

A cold front clipped northeastern Utah on April 29, with temperatures dropping to around -9° C at 700 mb. Moisture was limited, although westerly to northwesterly winds behind the front resulted in some snow showers during the morning and midday hours. Seeding was conducted at a couple of sites to target this activity. Precipitation totals were minimal, generally around 0.1" or less, and no icing activity was observed. Radiometer data showed some significant liquid water aloft (above 16,000 to 18,000 feet) on the night of April 28-29 although values below that appear very limited or uncertain.

Figure 4.10 shows precipitation as a percentage of the median values for the month of April.

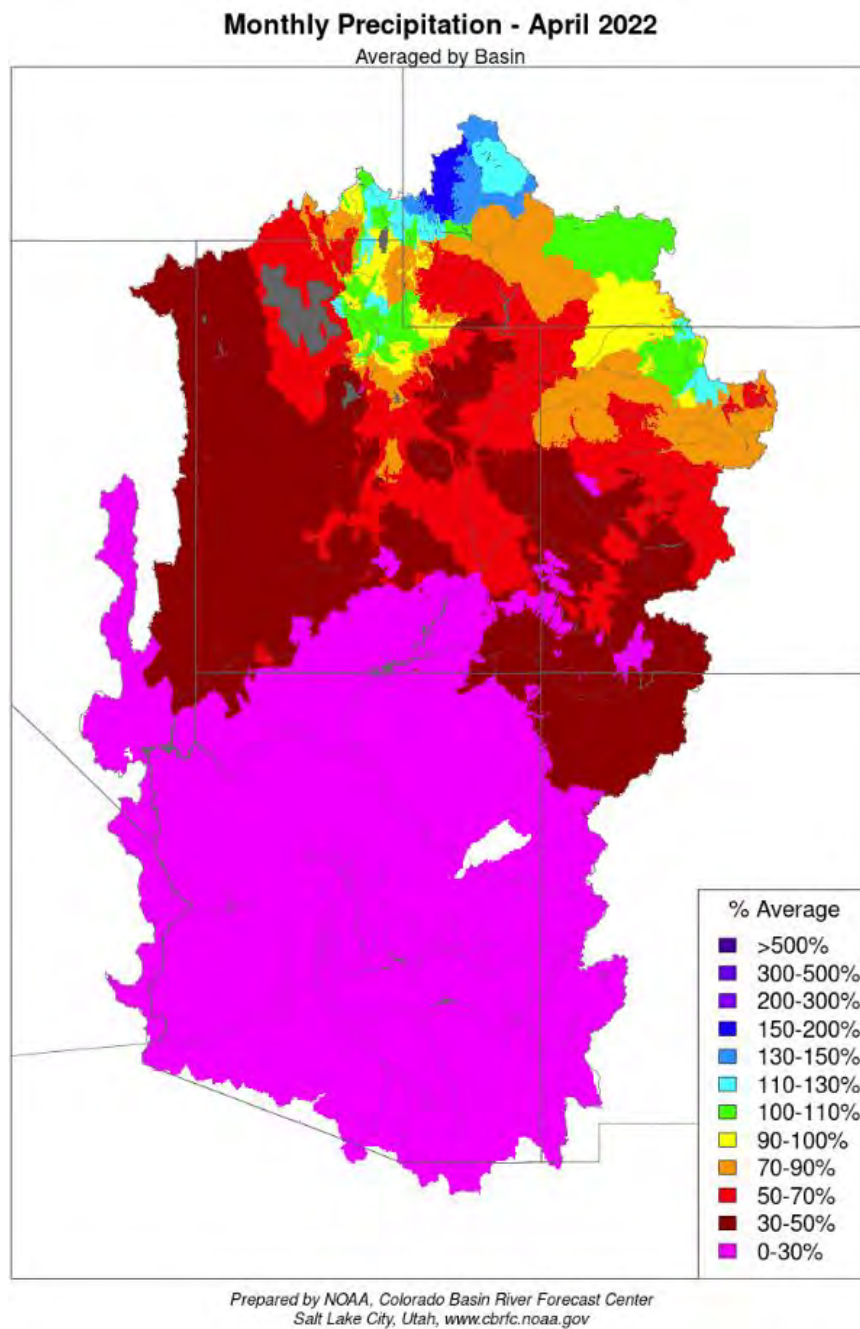


Figure 4.10 **April 2022 precipitation, percent of normal**

5. ASSESSMENT OF SEEDING EFFECTS

5.1 Background

The task of determining the effects of cloud seeding has received considerable attention over the years. Evaluating the results of a cloud seeding program for a particular season is rather difficult. The primary reason for the difficulty stems from the large natural variability in the amounts of precipitation that occur in a given area and between one area and another during a given season. The ability to detect a seeding effect becomes a function of the magnitude of the seeding increase and the number of seeded events, compared with the natural variability in the precipitation pattern. This can be described as a (seeding effect) signal to noise ratio issue. Larger seeding effects can be detected more easily, and with a smaller number of seeded cases, than are required to detect smaller increases.

Historically, consistently positive results have been observed in wintertime seeding programs in mountainous areas. However, the apparent differences due to seeding are relatively small, usually of the order of a 5-10 percent seasonal increase. In part, this relatively small percentage increase accounts for the significant number of seeded seasons (often ten years or more) required to establish these results for a particular program with any confidence.

Despite the difficulties involved, some techniques are available for estimation of the effects of operational seeding programs. These techniques are not as statistically rigorous as the randomization technique used in research, where roughly half the sample of storm events is randomly left unseeded. However, most of NAWC's clients do not choose to cut the potential benefits of a cloud seeding project in half in order to better document the effects of the cloud seeding project. The less rigorous techniques can, however, potentially offer a reasonable indication of the long-term effects of seeding on operational programs.

A commonly employed technique, the one utilized by NAWC in this assessment and in evaluation of its other winter seeding projects, is a "target" and "control" comparison. This technique is described by Dr. Arnett Dennis (1980) in his book entitled "Weather Modification by Cloud Seeding". The technique is based on the selection of a variable that would be affected by seeding (such as precipitation or snowpack). Records of the variable to be tested are acquired for an historical period of many years' duration (20 years or more if possible). These records are partitioned into those located within the designated "target" area of the project and those from well-correlated "control" sites located well outside of the target area.

Ideally the control sites should be selected in an area meteorologically similar to the target, but one which would be unaffected by the project seeding (or seeding from other adjacent projects). The historical data in both the target and control areas are taken from past years that have **not been subject to cloud seeding activities**. These data are evaluated for the same seasonal period of time (months) as that when the seeding is to be, or has been, conducted. The target and control sets of data for the unseeded seasons are used to develop a mathematical model (typically a linear regression), which predicts the amount of target area natural precipitation, based on precipitation observed in the control area. This mathematical model is then used to analyze the selected variable in years where seeding **did not** occur in the control but **did** occur in the target areas. From the model and available data for the control area, we can predict what

would likely have transpired in the target area had no seeding occurred, then compare this to what actually happened in the designated target area. Consistent differences between these predicted and observed values may be attributed to cloud seeding effects, although with a low level of confidence until sufficient seeded season data is accumulated.

This target and control technique works well where a good historical correlation can be found between target and control area precipitation. Generally, the closer the target and control areas are geographically, and in terms of elevation, the higher the correlation will be. Control areas selected too close to the target, however, can be subject to contamination by the seeding activities. This can result in an underestimate of the seeding effect. For precipitation and snowpack assessments, a correlation coefficient (r) of 0.90 or better would be considered excellent. A correlation coefficient of 0.90 would indicate that over 80 percent of the variance (r^2) in the historical data set would be explained by the regression equation used to predict the variable (expected precipitation or snowpack) in the seeded years. An equation indicating perfect correlation would have an r value of 1.0.

Experience has shown that it is virtually impossible to provide a precise assessment of the effectiveness of cloud seeding based on a small number of seeded seasons. However, as the data sample size increases, it becomes possible to provide at least a reasonable estimate of seeding effectiveness.

5.2 Target vs. Control Evaluations – Precipitation and Snowpack Data

The Natural Resources Conservation Service (NRCS) collects data from a number of precipitation and snow measurement sites. Most of these sites have been converted to automated SNOTEL sites in the last 30 years, although manual snow course measurements are still conducted at some locations. NAWC has utilized monthly precipitation and snow data from a number of these sites for use in seeding program evaluations. The number of sites operated by agencies such as the NRCS, especially manual snow course sites, has been gradually reduced. Even some cooperative observer sites, which are managed by the National Weather Service, have been either discontinued or have become inactive. Therefore, the selection of target and control sites first involves examination of the period of record of data at a given location, and changes to the set of target or control sites are sometimes necessary in the event that measurements at a site are discontinued.

There are multiple cloud seeding programs conducted in the State of Utah. As a consequence, potential control areas that are truly unaffected by cloud seeding are somewhat limited in geographic area. This is complicated by the fact that the best correlated control sites are generally those closest to the target area. Many measurement sites in this part of the state, although not located within the boundaries of the intended area of effect of a seeding program, have been subjected to potential effects of numerous historical and current seeding programs. This renders such sites of questionable value for use as control sites. Studies of downwind seeding effects suggest that if we wish to consider any precipitation gauge sites downwind of the seeded area as control sites for the High Uintas project, they should be located at least 50-75 miles downwind of current or historic cloud seeding programs in Utah (or Idaho and Nevada) to avoid significant contamination.

Our normal approach in selecting control sites for a new project is to look for sites upwind or crosswind from the target area that will geographically bracket the intended target area. The reason for this approach is that some winter seasons are dominated by one upper-level wind (jet stream) pattern while other seasons are dominated by other flow patterns. The result of these differing weather patterns and storm tracks often results in heavier precipitation in one area versus the other. For example, a strong El Niño pattern may favor below normal precipitation in one region more than in another. Having control sites on either side of the target area relative to the generalized flow pattern can improve the prediction of target area precipitation under these variable upper air weather patterns.

An additional consideration in the selection of control sites for the development of an historical target/control relationship is one of data quality. A potential control or target site may be rejected due to poor data quality if the data significantly diverges over time from other sites in the area. SNOTEL sites, the type used in the evaluation of the High Uintas program, typically have reliable long-term records with external variables (such as terrain aspect and surrounding vegetation) carefully selected or maintained. The double mass plot is an engineering tool that will indicate any changes in relationships between two stations, and may be particularly useful if one or both stations have moved during their history. If a site exhibits either an abrupt change due to relocation, or long-term trends that differ substantially from other sites in the area, it may be excluded from further consideration.

There are some things to consider when dealing with the two types of precipitation observations typically available from mountainous areas in the west: standpipe storage precipitation gauges and snow pillows. There are some potential problems associated with each type of observation. With the advent of the SNOTEL data acquisition system in the late 1970's, access to precipitation and snowpack (water equivalent) data in mountainous locations became routine. Before the SNOTEL system was developed, these data had to be acquired by actually visiting the site to make measurements. This is still required at some sites. Figure 5.1 is a photo of an NRCS SNOTEL site, with labels to allow the reader a better understanding of the two types of observation systems. The vertical tube is the standpipe storage gauge, which is approximately 12" in diameter. The gauges are approximately 20' in height so that their sampling orifices remain above the snowpack surface. There are at least two types of potential problems associated with high elevation observations of the water equivalent of snowfall, as measured by standpipe precipitation storage gauges. The two areas of concern are clogging at the top of the standpipe storage gauge, and blow-by of snowflakes past the top of the standpipe gauge. Either situation would result in an underestimate of the actual precipitation that fell during such periods. In the fall, the storage gauge is charged with antifreeze, which melts the snow that falls to the bottom of the gauge. A pressure transducer records the weight of the solution. The weight of the antifreeze is subtracted from the total weight, giving the weight of the water, which is then converted into inches. Heavy, wet snow may accumulate around the top of the standpipe storage gauge, either reducing or stopping snow from falling into the standpipe and resulting in an underestimate of precipitation. Snow that falls with moderate to strong winds may blow past the top of the gauge, which can also result in an underestimate of precipitation. NRCS sites are normally located in small clearings in forested areas to help reduce the impacts of wind effects. Sites that are near or above timberline are more likely to be impacted by wind since properly sheltered sites may be difficult to find in these areas. The snow pillow, pictured on the pad at ground level in the foreground

of Figure 5.1, is filled with antifreeze. This system weighs the snowpack, providing time-resolved records of the snowpack water content. Snow pillows can also have difficulty in providing accurate measurements of snow water content, because of wind either adding or removing snow from the measurement site when snow conditions are favorable for drifting.

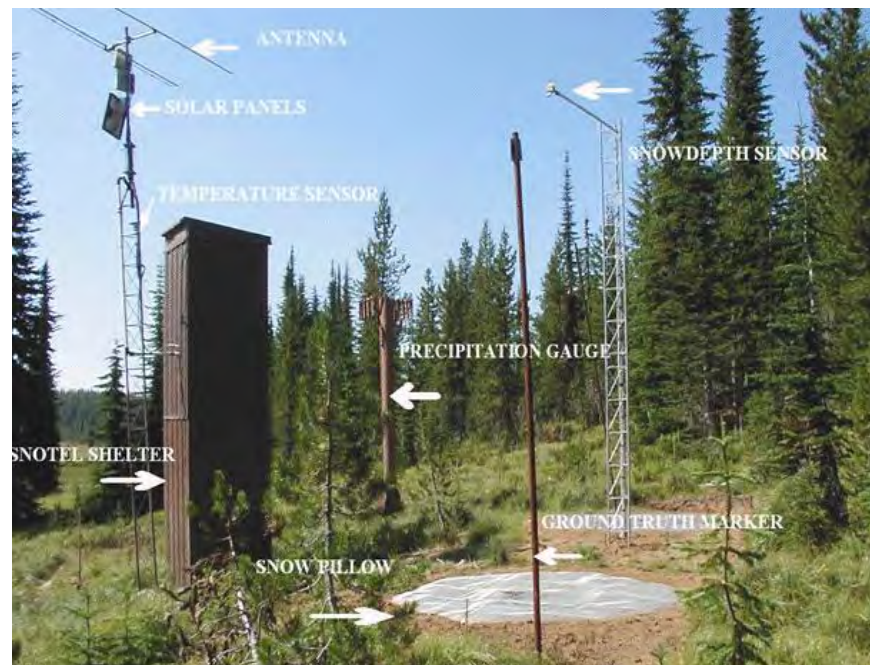


Figure 5.1 Equipment at a SNOTEL site

The water content within the snowpack is important since, after consideration of antecedent soil moisture conditions, it ultimately determines how much water will be available to replenish the supply when the snow melt occurs. Hydrologists routinely use snow water content measurements to forecast streamflow during the spring and early summer months. As with the precipitation storage gauge and SNOTEL precipitation gauge networks, Utah also has access to an excellent snow course and SNOTEL snow pillow reporting system via the NRCS. Many of the same reporting mountain sites are available for both precipitation and snowpack measurements. Consequently, it is worthwhile to evaluate the effects of seeding on snowpack as well.

There are some potential problems with snow course (manual) type of measurements that must be recognized when using those measurements to evaluate seeding effectiveness. Because not all winter storms are cold, sometimes rain as well as snow falls in the mountains. This can lead to a disparity between precipitation totals which theoretically measure everything that falls, and snowpack water content which measures only the water held in the snowpack. Warm periods can occur between snowstorms. If a significant warm period occurs, some of the precipitation that fell as snow will have melted or sublimated by the time the next snow course measurement is made. Thus, some of it may not be recorded in the snow water content measurements. This can also lead to a greater disparity between the snow water

content at higher elevations (where less snow will melt in warm weather) and that observed at lower elevations. The newer daily SNOTEL measurements avoid some of these problems, but depletion of the snowpack can occur even with SNOTEL measurements when dealing with April 1st observations. We are concerned with both types of measurements since we often use snow course measurements to provide a longer historical data base from which the regression equations can be developed. In addition, snowpack measurements are still conducted manually at a few mountain sites up to the present time.

Another factor that can affect the indicated results of the snowpack evaluation is the date on which snow course measurements were made. Since the advent of SNOTEL, data are now available on a daily basis. However, prior to SNOTEL, and at those sites where snow courses are still measured by visiting the site, the measurement is recorded on the day it was made. In some cases, because of scheduling issues or stormy weather, these measurements have been made as many as 5-10 days before or after the end of the month. This can lead to a disparity in the snowpack water content readings when comparing one group (such as a control) with another control or target group. Normally, however, snowpack measurements are made within a few days of the intended date. Nonetheless, the measurement timing issue can affect the data. Only two manual snow course sites are still used in analysis for this program, both of which are located in the target area.

April 1st snowpack readings have generally become accepted as the conventional data set for evaluating seasonal snowpack water content since they usually approximate the maximum snow accumulation for the winter season. Most streamflow and reservoir storage forecasts are made on the basis of the April 1st snowpack data. For that reason, and because five months of seeding are contained in the April 1st snowpack measurements, April 1st was selected as the most appropriate standardized date for snowpack analysis.

The bottom line is that it is difficult to accurately measure snow water equivalent at unmanned high-elevation sites. Both types of NRCS observations (gauge and snow pillow) can best be viewed as approximations of the actual amount of water that falls during a winter season. NRCS SNOTEL sites frequently provide the only type of precipitation observations available from the higher elevation areas targeted by winter cloud seeding programs. They are well-suited for use in estimations of seeding effects, but interpretation of the indicated seeding effects must take into account the limitations of the measurement systems and their data.

5.3 Target vs. Control Evaluations – Streamflow Data

In addition to the precipitation and snow water equivalent data which are used in these evaluations, NAWC has utilized some streamflow data for use in target and control analyses for the High Uintas program. Monthly streamflow data were obtained from the USGS (United States Geological Survey) website for sites that had a long history of unregulated streamflow measurements. Streamflow data can, under the right circumstances, directly address the issue of how much additional water is being produced by a seeding program. There are some potential difficulties here as well, including diversions for irrigation (which are present to some extent above even most of the “unregulated” sites), and significant carryover in soil moisture and resultant streamflow from one season to another, which lowers the correlation

between target and control sites. Overall, the best correlation between control and target sites is found with the precipitation data, followed by snow water equivalent, with streamflow correlations generally having the lowest correlation of the three data types.

5.4 Evaluation Methodology

Using the target-control approach introduced in the Section 5.1, the mathematical relationships for two variables (precipitation and snowpack) were determined between a group of sites in an unseeded area (the control group) and the sites in the seeded area (the target group), based upon records for a common period prior to any seeding in either area. From these data, mathematical models were developed whereby the amount of precipitation or snowpack observed in the unseeded (control) area was used to predict the amount of natural precipitation in the seeded (target) area. This “predicted” value is the amount of precipitation or snowpack that would be expected in the target area without seeding. The difference between the predicted amount and the observed amount in the target area is the excess, which may be the result of cloud seeding. Statistical tests have shown that such indications have little statistical significance for individual seasons, and usually fall within the standard deviation of the natural variability. However, more meaningful estimates can be obtained by combining the results of several or more seeded seasons.

5.5 Target and Control Sites - Precipitation

Precipitation measurements were available from six sites within the target area, the same sites as used in the previous several years. There are additional SNOTEL sites in the target area (e.g., Chepeta), but they have shorter periods of record. Thus, they were not considered in this analysis. The sites selected for use in the evaluation work are shown in Figure 5.2, and are all higher elevation NRCS sites. The average elevation for the target area sites is 9,875 feet above mean sea level (MSL). Specifics in regard to location and elevation of these six target area sites are provided in Table 5-1.

For many years, winter cloud seeding in Utah was limited to mainly the central and southern portions of the State, although occasional winter seeding was conducted in the mountains of Tooele County (southwest of the Salt Lake City area) in the late 1970’s and early 1980’s. However, beginning in the 1988 water year, winter cloud seeding programs became more widespread in northern Utah. The result of this increase in cloud seeding projects is that it has become more difficult to locate control areas that have not been affected by other cloud seeding programs. Also, some (non-SNOTEL) precipitation gage sites used as controls no longer have ongoing data collection.

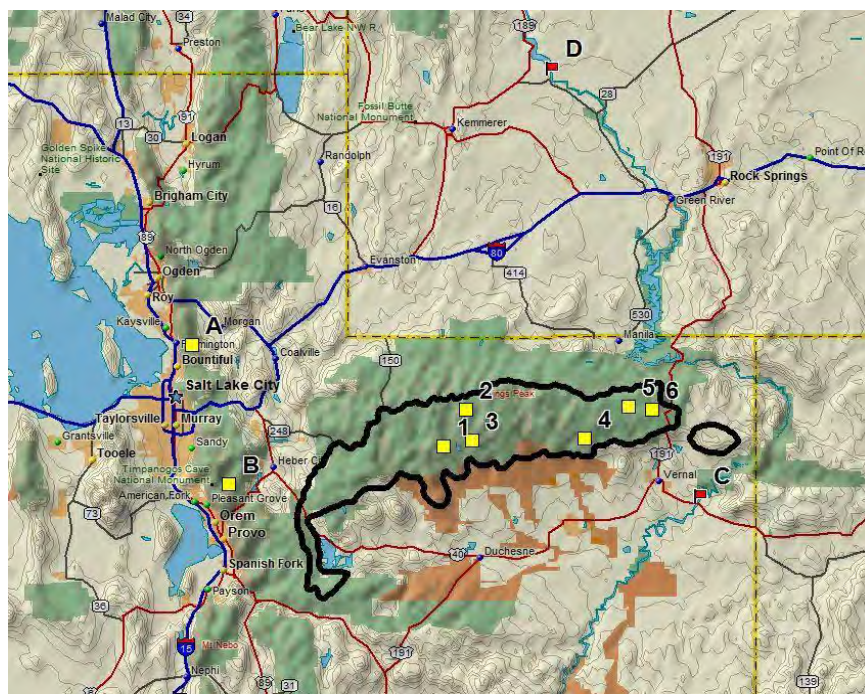


Figure 5.2 Precipitation gauges used as target area sites (number ID's) and control sites (letter ID's). The yellow boxes represent SNOTEL locations and the flag is an NWS co-op site.

The control gauge sites used in the evaluations were carefully selected according to the following criteria: 1) similarity to the target area sites, in terms of elevation and meteorology; 2) geographic bracketing of the target area; and 3) mathematical correlation of the data with that in the target area. The Strawberry Divide SNOTEL site was at one time included in the control group, but has been excluded from evaluations in recent years since it is now in part of the target area. Two cooperative (valley) reporting gauges, located at Heber and Vernal, were previously used as control sites, but have been discontinued because data are no longer available at these sites. The relationship of the control area gauges to the target area is shown in Figure 5.2, and the specifics in regard to the locations and elevations of the control sites are provided in Table 5-1.

**Table 5-1
Control and Target Area Precipitation Gauge Sites**

Group ID	Site Name	Site Number	Elevation (feet)	Latitude (N)	Longitude (W)
Control					
A	Farmington Canyon Upper	11J11S	8000	40°58'	111°48'
B	Timpanogos Divide	11J21S	8140	40°26'	111°37'
C	Jensen	424342	4750	40°22'	109°21'
D	Fontenelle Dam, WY	483396	6480	41°59'	110°04'
Target					
1	Brown Duck	10J30S	10600	40°35'	110°35'
2	Five Points Lake	10J26S	10920	40°43'	110°28'
3	Lakefork #1	10J10S	10100	40°36'	110°26'
4	Mosby Mountain	09J05S	9500	40°37'	109°53'
5	Trout Creek	09J16S	9400	40°44'	109°40'
6	King's Cabin	09J01S	8730	40°43'	109°33'

It is recognized that the group of control sites in Table 5-1 might provide a conservative estimate of the effects of seeding for the High Uintas, since there could have been some seeding effects impacting some of the control sites (e.g. seeding from the Western Uintas and Six Creeks programs could impact the precipitation at Heber, and seeding in eastern Tooele County and in Box Elder County could impact sites like Farmington Canyon). Those impacts would have the effect of raising the predicted target area precipitation and, thus, lowering the indicated effects of seeding in the High Uintas target area. The average elevation of all seven control sites is 6,842 feet, which is much lower than that of the target sites (9,875 feet). The large elevation difference is due in part to the fact that the Uinta Range is the highest mountain range in the region, and most other high elevation sites are in areas with other seeding programs. The locations of the control sites are shown in Figure 5.2. Elevation differences are important in snow water content evaluations because snowmelt may impact high and low elevation sites differently. The great elevation difference between the target and control sites is also of significance in the precipitation evaluations because of the potential for much windier exposures at the Uintas sites which are ~3,000 feet higher on average than the control sites. Gauge catch deficiency due to wind can be very significant in more exposed areas, as can the problem of drifting snow.

5.6 Target and Control Sites - Snowpack

The procedure was essentially the same as was done for the precipitation evaluation, e.g., control and target area sites were selected, and average values for each were determined from the historical snowpack data. Due to concerns regarding potential contamination by other seeding projects, combined with some period of record limitations and consideration of site correlation values, a short 13-year

historical period (1975-88) was used in most of the snow water content evaluations. The limited amount of historical data renders the equations using the historical regression technique questionable, as described in the earlier precipitation evaluation section. Historical periods of at least 20 seasons duration are very desirable when utilizing this technique. The years after the 1988 water year were excluded from the historical period in most of these evaluations, given a number of seeding programs in northern Utah beginning with the 1989 water year, especially along the Wasatch Range west of the Uintas.

Four sites were selected as controls for the snowpack evaluation. The control group provides reasonably good correlations with the six-site target area group. The six snowpack target sites include four of the six sites used in the precipitation evaluations (data were unavailable back to 1975 for the Brown Duck and Five Points Lake sites), plus two additional manual snow course sites (Lakefork Mountain #3 and Spirit Lake). Spirit Lake is actually located on the north slope of the Uintas but is very close to the crest, so we believe it to be representative of the original target area in general (and now within the target area, with the additional of the northern slope to the program). It should also be noted here that SNOTEL sites were installed in 2009 at the Lakefork Mountain #3 and Spirit Lake snow course locations, and data at these sites became SNOTEL-only (instead of snow course) beginning in 2011. The target and control area snow course/snow pillow site names, elevations and locations are summarized in Table 5-2, and site locations are shown in Fig. 5.3. The elevations of the control area sites averaged 8,184 feet. The target sites were significantly higher, averaging 9,405 feet. The relationship of the control area snowpack sites to the target area is shown in Figure 5.3.

Due to the challenges involved in the target/control analyses for the High Uintas program, including concern over short historical periods, a snow water content regression (linear and multiple linear) that uses fewer sites but a much longer historical regression period of 46 years was also conducted.

5.7 Target and Control Sites - Streamflow

NAWC has investigated numerous target/control type evaluation techniques, as well as multiple variations of existing techniques, in an attempt to provide the client with a reasonable estimate of precipitation increases resulting from the seeding program. One of these techniques is an evaluation based on March – July streamflow, utilizing several control sites that had essentially unregulated streamflow records. Three suitable control sites were located in western Wyoming, and two sites were similarly located in northwestern Colorado. Three suitable (unregulated) streamflow gauges were used to represent target area runoff (Yellowstone, Lake Fork and Ashley Creek drainages). Streamflow data at these sites have longer periods of record than SNOTEL snow and precipitation data, yielding a longer historical base period. The sites utilized in these streamflow comparisons have data back to at least 1964, allowing a 30 year base period to be established for the period prior to the beginning of the South Slope seeding program (certain years were excluded from the base period due to a historical seeding program affecting western Wyoming). There were two separate regions with unregulated streamflow gauges that were judged to be suitable for controls. One of these groups is in western Wyoming. Examination of the correlation between these and the target area sites, along with examination of double-mass plots, an engineering tool used to examine the consistency of an historical paired data set, resulted in three of these Wyoming gauges being selected as controls. Similarly, two control sites were selected from an available set in northwestern Colorado, which are unlikely to be affected by current or historical seeding programs. These sites are listed in Table 5-3, and shown on the map in Figure 5.4.

**Table 5-3 Control and Target Streamflow Gauges
(Data obtained from the USGS website)**

Group ID	Site Name	USGS Site Number	Latitude (N)	Longitude (W)
Control - Wyoming and Colorado				
A	Hams Fork, WY	09223000	42°07'	110°42'
B	Smith's Fork, WY	10032000	42°03'	110°24'
C	Fontenelle Creek, WY	09210500	42°06'	110°25'
D	Little Snake River, CO	09260000	40°33'	108°25'
E	White River	09304500	40°02'	107°51'
Target - Utah				
1	Lake Fork	09289500	40°36'	110°32'
2	Yellowstone River	09292500	40°31'	110°20'
3	Ashley Creek	09266500	40°35'	109°37'

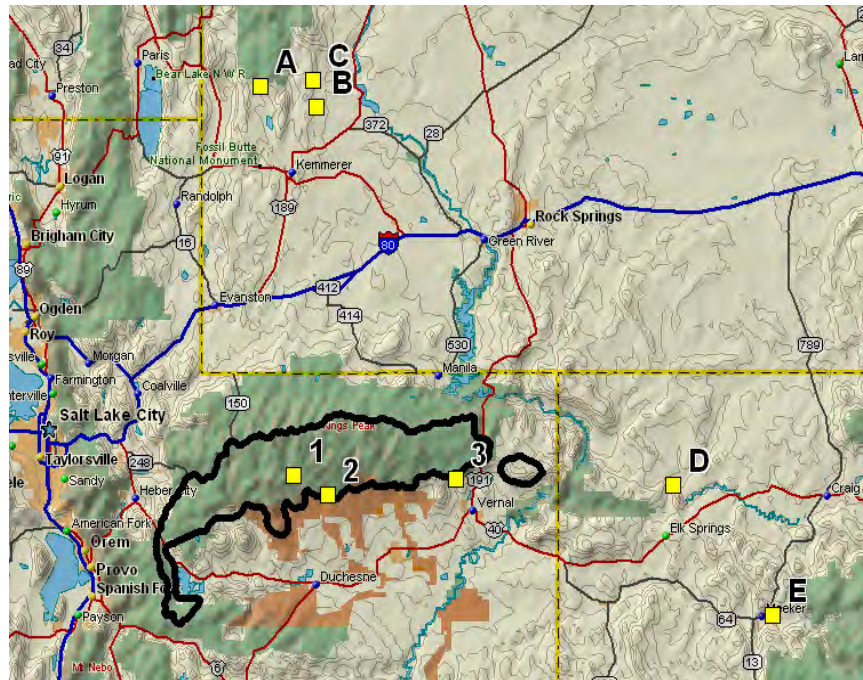


Figure 5.4 High Uintas streamflow target and control gauges

Over the course of this seeding program, several evaluation methods have been applied to the precipitation, snowpack and streamflow data. The results of the various evaluations are summarized in the following sub-sections, and Appendix C contains more detailed information for some of these evaluations.

5.8 Development of Regression Equations

NAWC compared various methods of analyzing the data, including the linear and multiple linear regression methods which have been used with this and similar programs. The target and control site historical (non-seeded) data for precipitation, snowpack, and streamflow were used to develop regression equations that describe the relationship between the control and target areas in the absence of cloud seeding. In the precipitation evaluation, for example, the monthly precipitation values were totaled at each gauge in the control and target areas for the December-April periods in each of the historical (not seeded) water years from 1980 - 1988, 1994, and 1996-2000, for a total of 15 seasons. The reasons for the short historical period are a) a lack of consistent precipitation measurements prior to the advent of the SNOTEL observations and b) the necessity of excluding winter seasons in which there were some seeding activities conducted in upwind areas that may have impacted precipitation in the High Uintas target area (e.g., projects in the western Uintas or the Wasatch Front area). Averages for each group were obtained, and predictor equations developed from these data for a five-month period (December through April). Appendix B contains details regarding some of the historical regression relationships that have been developed and applied to the seeded seasons.

Development of snowpack and streamflow regressions was similar. The snowpack analyses were based on snow water equivalent amounts measured on April 1st (using both the SNOTEL and snow course measurements). April 1st is important because it approximates the total seasonal snowpack accumulation fairly well in many areas, usually before significant melting begins. Also, many water supply forecasts are based on April 1 snow water content. The streamflow analysis utilized total streamflow (in acre-feet) during the March – July period. This period has been found to be one of the best correlated with winter season precipitation. April – July streamflow can be used for this as well, although the runoff can begin during March in some seasons, especially areas on a southerly exposure such as the southern slopes of the Uintas. The primary snowpack regression used for this program was based on only 13 historical seasons (water years 1975 – 1987), although an alternate snowpack regression that was also developed utilized long-term historical data available at only a small number of sites to produce a 46-year historical period. The streamflow regression was based on a fairly long historical period of 30 seasons. These include water years 1966, 1971-79, and 1983-2002. The historical regression periods were selected on the basis of data availability and avoidance of seasons where historical seeding programs would have directly impacted some or all of the control sites.

Multiple regression analyses relate each control site individually to the average of the target area sites, and these were conducted as well. This multiple regression analysis method was used because it provides a higher correlation between control and target sites, which can yield a better estimate of seeding effects if there is sufficient historical (non-seeded) data for a meaningful regression equation to be established using this method. For the precipitation and snowpack evaluations, a relatively short historical period makes this type of analysis somewhat questionable since the number of independent variables (control sites) in the equation becomes relatively large in comparison to seasons in the historical period. The results of the multiple regression analysis (for precipitation and snowpack) were still considered, but for this program the multiple regression method is better suited to the streamflow data set which has a much longer historical period.

5.9 Evaluation Results

Precipitation evaluation results have been examined for a period of 20 seeded seasons (2003-2022 water years). The seeded period used in one snowpack evaluation (with more sites but a short historical period) excludes the water year 2004, 2007, 2012, and 2015 seasons due to early melting in those years, and so includes only 15 seasons. The other long-term snowpack evaluation (few sites but 46 historical seasons) excludes these same seeded seasons due to early snow melt. This evaluation originally had three control sites but one snow course (White River #3) appears to have been discontinued in 2016 so the regression equation was re-established without this site. The streamflow evaluation currently has data available through 2021 for the March – July seasonal period, and so includes the 2003-2021 water years, for a total of 19 seasons.

The evaluation techniques as described yield an estimation of the observed/predicted amount of precipitation, snow water content, or streamflow for an individual season. Individual season results are included in the tables in Appendix B, in the “RATIO” column for the seeded seasons. Results for the current season are discussed below Table 5-4. A ratio of 1.05, for example, would suggest a 5% increase over the

natural precipitation, snowpack, or streamflow predicted for the target area based on the historical regression equation. A ratio at or below 1.0 is not indicative of an increase over the natural precipitation or snowfall. An increase for an individual seeded season or combination of seeded seasons could be attributed to seeding effects. However, it is important to exercise caution in interpreting single-season statistical indications, since the natural variability of weather patterns between control and target areas will often outweigh the effects of seeding in a given year. This natural variability can result in a false or exaggerated positive indication, or in a low ratio (lack of indicated effects) when seeded effects were actually present. The strength of this type of evaluation is in multi-season indications over many seeded years.

Table 5-4
Summary of High Uintas Evaluation Results

Evaluation Type	Method	Pre-Seeded Years	Seeded Years	Correlation (R-value)	Resultant Ratio
April 1 Snow Water Content	Linear Regression	13	16*	0.81	0.96
April 1 Snow Water Content	Multiple Linear	13	16*	0.94	1.04
April 1 Snow Water Content	Linear Regression	46	16*	0.83	1.03
April 1 Snow Water Content	Multiple Linear	46	16*	0.86	1.09
Dec – Apr Precipitation	Linear Regression	15	20	0.86	0.96
Dec – Apr Precipitation	Multiple Linear	15	20	0.92	0.95

* Snowpack result excludes 2004, 2007, 2012, and 2015 due to early snow melt

** Streamflow evaluation includes seeded year data up through 2021, as the full March – July streamflow data for the current season is not yet available

Overall, indications from the various evaluation methodologies (linear regression and multiple linear regression) were mixed. Appendix B contains detailed evaluation results. Overall, a majority of these observed/predicted ratios were in the 0.95 - 1.09 range, particularly for the evaluations that exhibit more stable mathematical characteristics (i.e. evaluations of December – April precipitation and snow water content). Correlation (expressed as R-values) was generally highest for the precipitation evaluations, somewhat lower for the snowpack evaluations, and lowest for the various streamflow evaluations. Relatively low correlations (R values of much less than perhaps 0.85) indicate that there is considerable natural variability between the control and target areas, which for the South Slope of the Uintas target area is essentially unavoidable given its uniqueness in terms of meteorology, climatology and barrier orientation. Development and performance of the regression equations are greatly affected by the duration of the historic period; longer base periods are highly desirable. Because of this factor, NAWC included a long-term snowpack evaluation, as mentioned earlier, using a base period of 46 seasons and a

limited number of target/control sites with long records, sites that are also unlikely to be affected by surrounding seeding programs. The results of this particular evaluation (ratios of 1.03 for the linear and 1.09 for the multiple linear, for the average of the seeded seasons) are suggestive of snowpack increases during the seeded seasons for the High Uintas seeding program. Snowpack evaluations were not meaningful for the 2004, 2007, 2012 and 2015 seasons due to substantial early snowmelt and those seasons were excluded from the snowpack evaluation results.

With regard to the streamflow data, it is apparent in the data that not only are the correlations fairly poor, but there appears to be a good deal of non-linearity that produces better evaluations results (in general) during wetter years. This is likely due to higher inter-annual variability in the target area than the control areas. It is also observed that the seeded seasons were generally drier as a whole, on the order of 10% less runoff, than the historical seasons across the entire region (both target and control). Thus, the seeded season results have likely been biased toward lower ratios given the overall dryness of these seasons.

It is important to recall that, for the High Uintas program, there are a number of factors that make a meaningful analysis of the seeding effects difficult. These include the following: a) a relatively small number of seeded seasons, b) high seasonal variability between control and target areas, c) generally short historical periods without seeding from which regression equations can be developed, d) potential impacts on the historical regression equations from other NAWC winter seeding programs, e) sensitivity to early snowmelt issues at south-slope locations, and f) the possible long-term reduction of precipitation in the target area due to pollution as documented for precipitation sites slightly west of the High Uintas target area (Griffith et al., 2005). Items b) and d) above are described more fully in sections below.

5.10 Seasonal Variability, Related to Storm Track and Barrier Orientation (item b)

From a meteorological standpoint, there are several possible reasons why target area precipitation was comparatively low on average during the seeding seasons compared to that observed in various control areas. The El Nino/La Nina phase and various other factors can affect the location and orientation of the primary storm track on a seasonal and multi-seasonal basis. This can lead to large (either negative or positive) precipitation anomalies in the High Uintas in comparison to the surrounding region, especially given the east-west orientation of the mountain barrier. Storm events that are accompanied by a wind pattern moving essentially straight west to east, i.e., basically barrier-parallel can present reasonable seeding opportunity for the target area, although base (natural) amount of precipitation falling in the High Uintas with this type of flow pattern is low compared to surrounding areas. The predominantly north-south oriented mountain barriers in the intermountain region produce strong orographic (terrain-induced) lift in westerly air flow situations, while the west-east oriented Uinta Range produces minimal lift in those situations. The result is a minimal orographic component of the precipitation in the Uintas during periods of westerly flow. Given that the orographic component of precipitation is high in the mountains of Utah, approaching 75% of the winter precipitation in many areas, a dominant wind pattern that is even slightly anomalous can lead to a negative precipitation anomaly that may more than offset the actual seeding effects. In addition, there are indications that large, closed-circulation storm systems (so-called cutoff lows) during the spring, which climatologically contribute a substantial amount of

snowfall over the Uinta Range particularly during the month of April, were relatively lacking during many of the seeded seasons. The effect of that sort of natural variation can easily mask positive seeding effects.

5.11 Impacts from Other Seeding Projects (item d)

Other seeding programs being conducted in Utah may be impacting the apparent effects of seeding in the High Uintas. For example, the programs conducted in Tooele County and Box Elder County (which included seeding in both western and eastern portions of the county last winter) may be increasing the precipitation at some of the northern control sites (e.g., Farmington Canyon) and seeding in Juab and Sanpete Counties could be increasing precipitation at some of the southern control sites (e.g., Timpanogos Divide and Heber). Some of the Uinta program SNOTEL sites are within approximately 50 miles downwind of other seeding programs. Solak et al. (2003) reported that precipitation appears to have been increased at similar downwind distances due to the cloud seeding program being conducted in central and southern Utah, with similar results in a subsequent analysis up through 2018. For the High Uintas precipitation evaluation, 15 historical seasons were selected which exclude Water Years 1989 through 2002 since a number of seeding programs began in WY 1988 or 1989 in northern Utah, especially along the Wasatch Range west (upwind) of the Uintas. These seasons were excluded from the historical period due to potential contamination effects. Similar exclusions resulted in a 13-year historical data set for the snowpack evaluation, while the streamflow evaluation had a different set of historical seasons (during the 1970s and early 1980s) excluded because of the Bear River seeding program affecting portions of western Wyoming where some of the streamflow control sites are located.

In order to illustrate the potential effects of contamination, assume that the average precipitation at the control sites was increased by 5% due to seeding from other programs. This would also raise the predicted target area precipitation by roughly 5%. If this were the case, it would cause a similar 5% precipitation increase in the High Uintas target area to be undetected in a more basic mathematical analysis. A final (and very important) consideration in the estimation of seeding effects for this program pertains to the results obtained from numerous similar programs in Utah and elsewhere in the western U.S. While each program is unique, evaluation results from most of these programs have ranged from approximately 5-10% increases over the estimated natural seasonal precipitation.

5.12 The Bottom Line

With consideration given to the meteorology and physiography of the Uintas, the range of results of various evaluations of seeding effects, the peculiarities of the seeded period, and results of similar programs, our best estimate is that the High Uintas seeding program has increased the project target area precipitation by approximately 3-5% on average during the seeded seasons. Table 5-4 summarizes the results of the various evaluations conducted to date for the High Uintas program. Detailed data from these evaluations are shown in Appendix C.

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APPENDIX A: SUSPENSION CRITERIA

Certain situations require temporary or longer-term suspension of cloud seeding activities, with reference to well-considered criteria for consideration of possible suspensions, to minimize either an actual or apparent contribution of seeding to a potentially hazardous situation. The ability to forecast (anticipate) and judiciously avoid hazardous conditions is very important in limiting any potential liability associated with weather modification and to maintain a positive public image.

There are three primary hazardous situations around which suspension criteria have been developed. These are:

1. Excess snowpack accumulation
2. Rain-induced winter flooding
3. Severe weather

Excess Snowpack Accumulation

Snowpack begins to accumulate in the mountainous areas of Utah in November and continues through April. The heaviest average accumulations normally occur from January through March. Excessive snowpack water content becomes a potential hazard during the resultant snowmelt. The Natural Resources Conservation Service (NRCS) maintains a network of high elevation snowpack measurement sites in the State of Utah, known as the SNOTEL network. SNOTEL automated observations are now readily available, updated as often as hourly. The following set of criteria, based upon observations from these SNOTEL site observations, has been developed as a guide for potential suspension of operations.

Project & Basin	Critical Streamflow Volume (Acft) & USGS Streamgage	SNOTEL Station	SWE Value Corresponding to the Critical Flow								Ranking of SNOTEL Stations
			Jan 1 (in.)	Jan 1 (%)	Feb 1 (in.)	Feb 1 (%)	March 1 (in.)	March 1 (%)	April 1 (in.)	April 1 (%)	
1. Northern Utah	183,708	Franklin Basin, Idaho	19.50	190.84	27.14	165.31	34.35	154.71	41.56	153.60	1
Logan at Logan	USGS 10109000	Tony Grove	28.73	205.94	39.44	175.56	48.06	160.38	56.34	156.56	2
		Bus Lake	17.08	218.82	21.91	180.34	26.72	165.25	31.65	162.70	3
		Average	21.80	205.20	29.50	173.70	36.40	160.10	43.20	157.60	
Weber near Oakley	USGS 10128500	Chalk Creek #1	10.09	173.13	14.73	153.66	28.77	149.85	34.15	143.41	1
		Trial Lake	20.15	207.44	26.33	180.55	33.55	173.27	38.54	162.28	2
		Smith Morehouse	10.00	186.34	13.89	177.60	17.36	146.22	21.17	160.26	3
		Hayden Fork	12.19	194.16	16.69	172.11	20.71	158.56	21.79	164.64	4
		Average	13.10	190.30	17.90	166.00	25.10	157.10	28.90	157.70	
Dunn Creek near the Park Valley	USGS 10172952	George Creek	17.84	187.75	18.32	143.81	28.93	163.43	34.61	153.77	1
		Howell Canyon, Idaho	28.71	279.96	38	223.24	44.59	205.98	50.46	191.65	2
		Average	23.30	233.90	28.20	183.60	36.80	184.70	42.60	172.70	
2. Western & High Uintah	166,861	Lily Lake	11.38	202.70	16.40	194.06	17.69	147.37	28.93	139.19	1
Bear River near Utah - Wyoming state line	USGS 10011500	Trial Lake	20.07	206.54	26.56	182.26	33.68	173.94	38.49	162.05	2
		Hayden Fork	12.41	197.65	17.08	175.83	21.03	160.98	20.90	146.02	3
		Average	14.60	202.30	20.00	184.10	24.10	166.80	29.40	149.10	
Duchesne near Tablona	USGS 09277500	Strawberry Divide	6.92	239.23	10.87	199.25	26.77	178.78	29.75	179.05	1
		Daniels-strawberry	16.07	248.12	21.50	202.44	27.82	190.54	29.89	182.75	2
		Smith Morehouse	10.61	196.64	14.95	172.41	17.82	158.83	22.22	168.26	3
		Rock Creek	8.76	230.02	12.31	219.65	15.88	205.68	16.41	209.06	4
		Average	10.60	228.50	14.90	198.50	22.30	183.50	24.60	187.30	
Provo near woodland	USGS 09277500	Trial Lake	22.98	236.53	27.78	190.63	35.23	181.59	31.44	132.39	1
		Beaver Divide	10.29	210.39	14.11	179.49	17.45	170.83	20.18	200.3	2
		Average	16.70	223.50	20.90	185.10	26.30	176.20	25.80	166.40	
3. Central & Southern	120,473	Castle Valley	12.23	244.05	16.96	203.04	22.22	187.68	26.30	180.00	1
Sevier near Hatch	USGS 10174500	Harris Flat	8.71	298.76	15.25	273.59	24.16	222.09	21.15	209.77	2
		Farnsworth Lake	17.35	218.10	20.66	185.95	27.05	182.74	32.93	167.03	3
		Average	12.80	253.70	17.70	220.90	24.50	197.70	26.80	185.60	
Coal Creek near Cedar City	USGS 100242000	Midway Valley	20.89	215.65	29.12	194.04	35.89	176.99	42.29	167.97	1
		Webster Flat	13.57	232.46	18.70	197.95	24.30	184.64	24.93	181.12	2
		Average	17.20	224.10	23.90	196.00	30.10	180.90	33.60	174.60	
South Willow near Grantsville	USGS 100172800	Rocky Basin-settlement	19.09	205.33	23.75	174.14	32.11	171.39	40.01	167.51	1
		Mining Fork	16.31	243.66	20.74	177.04	27.81	171.79	32.19	168.74	2
		Average	17.70	224.50	22.30	178.60	30.00	171.60	36.10	168.10	
Virgin River at Virgin	USGS 09406000	Kolob	23.11	229.25	29.08	220.78	36.51	197.43	43.71	196.21	1
		Harris Flat	9.71	377.00	15.69	304.18	21.46	300.00	20.11	370.00	2
		Midway Valley	24.76	256.17	34.56	238.40	41.44	209.68	51.05	211.06	3
		Long Flat	9.38	265.88	13.54	286.16	19.20	286.18	18.91	187.00	4
		Average	16.70	282.10	23.20	262.40	29.70	248.40	33.40	241.10	
Santa Clara above Baker Reservoir	USGS 09409100	Gardner Peak	13.00	293.90	16.82	172.15	21.70	167.36	24.45	163.95	1
		Average	13.00	293.90	16.80	172.10	21.70	167.40	24.50	164.00	
		Utah State Average (%)	230		197		183		178		
	Standard Deviation	42		38		35		42			
	Upper 95%	248		213		199		196			
	Lower 95%	212		180		168		160			

Snowpack-related suspension considerations will be assessed on a geographical division or sub-division basis. The NRCS has divided the State of Utah into 13 such divisions as follows: Bear River, Weber-Ogden Rivers, Provo River-Utah Lake-Jordan River, Tooele Valley-Vernon Creek, Green River, Duchesne River, Price-San Rafael, Dirty Devil, South Eastern Utah, Sevier River, Beaver River, Escalante River, and Virgin River. The Weber-Ogden and Provo River – Utah Lake – Jordan River criteria apply to suspension considerations for the Western Uintas project. Since SNOTEL observations are available on a daily basis, suspensions (and cancellation of suspensions) can be made on a daily basis using linear interpolation of the first of month criteria. For the High Uintas Program, SNOTEL sites including Lily Lake, Trial Lake, Hayden Fork, Strawberry Divide, Daniels-Strawberry, and Rock Creek have date-specific snow water equivalent criteria on which suspension decisions can be based.

Streamflow forecasts, reservoir storage levels, soil moisture content and amounts of precipitation in prior seasons are other factors which need to be considered when the potential for suspending seeding operations due to excess snowpack water content exists.

Rain-induced Winter Floods

The potential for wintertime flooding from rainfall on low elevation snowpack is fairly high in some (especially the more southern) target areas during the late winter/early spring period. Every precaution must be taken to insure accurate forecasting and timely suspension of operations during these potential flood-producing situations. The objective of suspension under these conditions is to eliminate both the real and/or perceived impact of weather modification when any increase in precipitation has the potential of creating a flood hazard.

Severe Weather

During periods of hazardous weather associated with both winter orographic and convective precipitation systems it is sometimes necessary or advisable for the National Weather Service (NWS) to issue special weather bulletins advising the public of the weather phenomena and the attendant hazards. Each phenomenon is described in terms of criteria used by the NWS in issuing special weather bulletins. Those that may be relevant in the conduct of winter cloud seeding programs include the following:

Winter Storm Warning - This is issued by the NWS when it expects heavy snow warning criteria to be met, along with strong winds/wind chill or freezing precipitation.

Flash Flood Warning - This is issued by the NWS when flash flooding is imminent or in progress. In the Intermountain West, these warnings are generally issued relative to, but are not limited to, fall or spring convective systems.

Seeding operations may be suspended whenever the NWS issues a weather warning for or adjacent to any target area. Since the objective of the cloud seeding program is to increase winter snowfall in the mountainous areas of the state, operations will typically not be suspended when Winter Storm Warnings are issued, unless there are special considerations (e.g., a heavy storm that impacts Christmas Eve travel).

Flash Flood Warnings are usually issued when intense convective activity causing heavy rainfall is expected or is occurring. Although the probability of this situation occurring during our core operational seeding periods is low, the potential does exist, especially over southern sections of the state during late March and early April, which can include the project spring extension period. The type of storm that may cause problems is one that has the potential of producing 1-2 inches (or greater) of rainfall in approximately a 24-hour period, combined with high freezing levels (e.g., > 8,000 feet MSL). Seeding operations will be suspended for the duration of the warning period in the affected areas.

NAWC's project meteorologists have the authority to temporarily suspend localized seeding operations due to development of hazardous severe weather conditions even if the NWS has not issued a warning. This would be a rare event, but it is important for the operator to have this latitude.

APPENDIX B: SEEDING OPERATIONS TABLES, 2021-2022

Table B-1
Generator Hours for High Uintas Program, 2021-2022, Storms 1-10 (rounded to quarter hour)

Storm	1*	2*	3	4	5	6	7	8	9	10
Dates	Nov 9-10	Nov 19-20	Dec 8-9	Dec 14-15	Dec 24-25	Dec 26	Dec 30-31	Jan 5-6	Jan 7-8	Feb 16
SITES										
HU1			4.75				16			
HU2			15.75				16		13	
HU3				12						10.5
HU4		1						21.25		12
HU5	7		6	9	16	6	11.75			
HU6	7		6	9	16	6	11.75			
HU7	16.5		6	11.25	23.25					
HU8	16.25		6	11	23					
HU9**				6		10.75				
HU10										10.5
HU11										6.25
HU12							12			
HU13	16			11.5	18.5		12			
HU14							12			
HU15										
HU16										
HU17										
HU18										10.75
HU19										8.5
HU20										8.5
HU22										
WU3										
WU4										
WU6	13									
WU7	12.5									
WU8										
WU9	13							22.75		
WU10		12.75				6.5		22	13	
WU11									13	
WU12				10		6.5				
Storm Total	101.25	13.75	44.5	73.75 + 6	96.75	25 + 10.75	91.5	66	39	67

*Seeding for Lower Basin Extension

** HU9 is a remotely operated site with higher output, hours are counted separately

Table B-2
Generator Hours for High Uintas Program, 2021-2022, Storms 11-20

Storm	11	12	13	14	15	16	17	18	19	20
Dates	Feb 21	Mar 4	Mar 5-6	Mar 8-9	Mar 13	Mar 16	Mar 20	Mar 29	Apr 4-5	Apr 11-12
SITES										
HU1										9
HU2					8.75		4		15	24.75
HU3								6.5		
HU4			17.25			4.75		10.75	16	24.5
HU5		5.25	9							
HU6		5.25	9							
HU7		8.75	12.5						13	19
HU8		9	12						10	19
HU9**		2	3							
HU10								10.75		
HU11								7.25		
HU12		4.75								
HU13			21.5							
HU14		4.5	8.75							
HU15	2.75	5	7							
HU16		5	7							
HU17										
HU18	4.5							9.5		
HU19	6.5		16					10.25		
HU20	6.25		16							
HU22			3							
WU3									16	14
WU4									16	24.75
WU5									16	20
WU6									16	20
WU7									17	9
WU8									16.5	16.25
WU9					4.75			4.5	15	
WU10				24.25	5.75			3	15	25.75
WU11								8.75	16.5	24.75
WU12				24.5			8.75	9.5	14	24.25
Storm Total	20	47.5 + 2	139 + 3	48.75	19.25	4.75	12.75	80.75	212	275

** HU9 is a remotely operated site with higher output, hours are counted separately

Table B-3
Generator Hours for High Uintas Program, 2021-2022, Storms 21-23

Storm	21	22	23	
Dates	Apr 14-15	Apr 22-23	Apr 29	Site Totals
SITES				
HU1				29.75
HU2		14		111.25
HU3		21	4.5	54.5
HU4		14		121.5
HU5	15	7		92
HU6	15	7		92
HU7	15.5			125.75
HU8	15.5			121.75
HU9				21.75
HU10				21.25
HU11				13.5
HU12		5.75		22.5
HU13				79.5
HU14		6		31.25
HU15		5.5		20.25
HU16		6		18
HU17				0
HU18				24.75
HU19				41.25
HU20				30.75
HU22		4		7
WU3				30
WU4		20		60.75
WU5				36
WU6				49
WU7				38.5
WU8				32.75
WU9		24.5		84.5
WU10		21		149
WU11		11		74
WU12		19.5	2.5	119.5
Storm	61	186.25	7	

APPENDIX C: EVALUATION DATA

Summary of High Uintas Evaluation Results

Evaluation Type	Method	Pre-Seeded Years	Seeded Years	Correlation (R-value)	Resultant Ratio
Dec – Apr Precipitation	Linear Regression	15	20	0.86	0.96
Dec – Apr Precipitation	Multiple Linear	15	20	0.92	0.95
April 1 Snow Water Content	Linear Regression	13	16*	0.81	0.96
April 1 Snow Water Content	Multiple Linear	13	16*	0.94	1.04
April 1 Snow Water Content	Linear Regression	46	16*	0.83	1.03
April 1 Snow Water Content	Multiple Linear	46	16*	0.86	1.09
March – July Streamflow 5 control 3 target	Linear Regression	30	19**	0.75	0.99
March – July Streamflow 5 control 3 target	Multiple Linear	30	19**	0.79	0.95
March – July Streamflow 3 control 3 target	Linear Regression	30	19**	0.61	0.95
March – July Streamflow 3 control 3 target	Multiple Linear	30	19**	0.63	0.93

* Snowpack result excludes 2004, 2007, 2012, and 2015 due to early snow melt

** Streamflow evaluation includes seeded year data up through 2021, as the full March – July streamflow data for the current season is not yet available

APPENDIX D: DETAILED EVALUATION DATA AND RESULTS

High Uintas December – April Precipitation, Linear Regression

Regression (non-seeded) period:

<u>Water Year</u>	<u>Control Avg</u>	<u>Target Avg</u>
1980	18.72	17.28
1981	11.03	9.75
1982	21.05	15.50
1983	16.37	13.12
1984	16.62	11.72
1985	10.70	11.50
1986	19.81	16.13
1987	7.85	9.78
1988	8.81	9.33
1994	12.22	10.95
1996	16.21	14.15
1997	18.09	16.83
1998	17.68	14.43
1999	14.03	15.32
2000	13.93	13.63
Mean	14.87	13.30

Seeded period:

<u>Water Year</u>	<u>Control Avg</u>	<u>Target Avg</u>	<u>Predicted</u>	<u>Ratio</u>	<u>Increase</u>
1989*	12.17	11.05	11.77	0.94	-0.72
1990*	10.68	13.47	10.92	1.23	2.54
1991*	12.21	11.62	11.79	0.99	-0.17
1992*	6.25	7.15	8.42	0.85	-1.27
1993*	15.77	16.45	13.80	1.19	2.65
1995*	15.80	15.15	13.82	1.10	1.33
2001*	12.27	13.93	11.83	1.18	2.11
2002*	11.15	7.83	11.19	0.70	-3.36
2003	9.32	9.40	10.16	0.93	-0.76
2004	13.84	12.15	12.71	0.96	-0.56
2005	18.91	17.20	15.57	1.10	1.63
2006	19.23	14.73	15.76	0.93	-1.02
2007	9.42	8.45	10.22	0.83	-1.77
2008	15.29	13.22	13.53	0.98	-0.31
2009	17.46	13.67	14.76	0.93	-1.09
2010	13.15	12.08	12.32	0.98	-0.24
2011	21.95	17.23	17.29	1.00	-0.06
2012	9.48	8.23	10.25	0.80	-2.02
2013	9.84	10.68	10.45	1.02	0.23
<u>Water Year</u>	<u>Control Avg</u>	<u>Target Avg</u>	<u>Predicted</u>	<u>Ratio</u>	<u>Increase</u>
2014	11.57	9.83	11.43	0.86	-1.60

2015	8.56	7.20	9.73	0.74	-2.53
2016	14.27	12.27	12.95	0.95	-0.69
2017	23.26	20.63	18.03	1.14	2.60
2019	19.35	16.17	15.82	1.02	0.35
2020	11.30	10.58	11.28	0.94	-0.69
2021	10.11	9.62	10.60	0.91	-0.99
2022	10.11	11.25	10.61	1.06	0.64
Seeded Mean	13.80	12.16	12.69	0.96	-0.54

* Seeding conducted in nearby areas but not in target area

SUMMARY OUTPUT

<i>Regression Statistics</i>	
Multiple R	0.858476
R Square	0.736981
Adjusted R Square	0.716749
Standard Error	1.417657
Observations	15

ANOVA

	<i>df</i>	<i>SS</i>	<i>MS</i>	<i>F</i>	<i>Significance F</i>
Regression	1	73.20731	73.20731	36.42607	4.2E-05
Residual	13	26.12676	2.009751		
Total	14	99.33406			

	<i>Coefficients</i>	<i>Standard Error</i>	<i>t Stat</i>	<i>P-value</i>	<i>Lower 95%</i>
Intercept	4.895582	1.439077	3.40189	0.004725	1.786645
X Variable 1	0.564797	0.093581	6.035401	4.2E-05	0.362628

High Uintas December – April Precipitation, Multiple Linear Regression

Regression (non-seeded) period:

<u>Water Yr</u>	<u>Timp Div</u>	<u>Farm Cyn</u>	<u>Jensen</u>	<u>Font Dam</u>	<u>Target avg</u>
1980	30.4	37.9	4.0	2.6	17.3
1981	18.3	21.0	2.8	2.1	9.8
1982	34.6	45.3	2.5	1.8	15.5
1983	22.5	36.6	3.5	2.8	13.1
1984	20.6	40.8	2.5	2.6	11.7
1985	18.9	19.6	3.4	0.9	11.5
1986	30.5	41.9	3.8	3.1	16.1
1987	10.6	16.8	2.4	1.6	9.8
1988	11.8	18.8	3.2	1.4	9.3
1994	18.8	27.2	1.7	1.2	11.0
1996	24.6	35.9	2.3	2.0	14.2
1997	28.0	37.6	4.0	2.7	16.8
1998	24.8	39.3	3.6	3.1	14.4
1999	18.9	30.1	3.8	3.4	15.3
2000	20.4	31.2	2.9	1.3	13.6
Mean	22.2	32.0	3.1	2.2	13.3

Seeded period:

<u>Water Year</u>	<u>Timp Div</u>	<u>Farm Cyn</u>	<u>Jensen</u>	<u>Font Dam</u>	<u>Target avg</u>	<u>Predicted</u>	<u>Ratio</u>	<u>Increase</u>
1989*	17.7	28.5	1.6	0.9	11.1	10.1	1.10	1.0
1990*	20.8	18.3	2.6	1.0	13.5	11.4	1.18	2.0
1991*	17.2	26.7	3.2	1.7	11.6	12.0	0.97	-0.4
1992*	9.2	13.0	1.8	1.0	7.2	7.6	0.94	-0.5
1993*	25.3	29.9	5.7	2.2	16.5	17.0	0.97	-0.6
1995*	25.3	32.2	2.9	2.8	15.2	13.9	1.09	1.2
2001*	16.9	28.1	2.1	2.0	13.9	10.7	1.30	3.2
2002*	13.3	28.2	1.2	1.9	7.8	8.7	0.90	-0.9
2003	11.0	21.8	2.7	1.8	9.4	9.7	0.97	-0.3
2004	17.6	32.0	2.3	3.4	12.2	11.6	1.05	0.6
2005	33.1	34.4	4.0	4.1	17.2	17.3	0.99	-0.1
2006	29.3	43.6	2.2	1.8	14.7	14.4	1.02	0.4
2007	12.8	20.8	2.8	1.3	8.5	10.1	0.83	-1.7
2008	21.4	33.5	4.6	1.6	13.2	15.0	0.88	-1.8
2009	25.7	38.1	4.4	1.7	13.7	15.9	0.86	-2.2
2010	21.5	25.0	3.9	2.2	12.1	13.8	0.88	-1.7
2011	36.0	45.5	4.4	1.8	17.2	18.7	0.92	-1.4

<u>Water Year</u>	<u>Timp Div</u>	<u>Farm Cyn</u>	<u>Jensen</u>	<u>Font Dam</u>	<u>Target avg</u>	<u>Predicted</u>	<u>Ratio</u>	<u>Increase</u>
2012	16.1	20.0	1.2	0.7	8.2	8.7	0.95	-0.5
2013	12.4	22.7	3.2	1.1	10.7	10.5	1.02	0.2
2014	16.3	25.6	2.5	1.8	9.8	10.9	0.90	-1.1
2015	11.4	19.9	1.7	1.3	7.2	8.4	0.86	-1.2
2016	20.4	30.8	2.9	3.0	12.3	12.7	0.96	-0.5
2017	37.9	44.5	3.8	6.8	20.6	19.2	1.07	1.4
2018	15.6	20.9	1.2	1.0	8.5	8.8	0.97	-0.2
2019	31.0	37.8	4.9	3.6	16.2	18.1	0.89	-1.9
2020	14.9	24.5	3.1	2.8	10.6	11.3	0.93	-0.8
2021	14.4	22.3	2.5	1.3	9.6	10.1	0.95	-0.5
2022	15.4	20.9	2.9	1.3	11.3	10.8	1.05	0.5
Seeded Mean	20.7	29.2	3.1	2.2	12.2	12.8	0.95	-0.6

* Seeding conducted in nearby areas but not in target area

SUMMARY OUTPUT

<i>Regression Statistics</i>	
Multiple R	0.92059
R Square	0.84749
Adjusted R Square	0.78649
Standard Error	1.23083
Observations	15

ANOVA

	<i>df</i>	<i>SS</i>	<i>MS</i>	<i>F</i>	<i>Significance F</i>
Regression	4	84.18464	21.046	13.8924	0.0004
Residual	10	15.14942	1.5149		
Total	14	99.33406			

	<i>Coefficients</i>	<i>Standard Error</i>	<i>t Stat</i>	<i>P-value</i>	<i>Lower 95%</i>	<i>Upper 95%</i>	<i>Lower 95.0%</i>	<i>Upper 95.0%</i>
Intercept	2.50414	1.804771	1.3875	0.19543	-1.5171	6.5254	-1.517	6.525418
X Variable 1	0.22402	0.122163	1.8338	0.09658	-0.0482	0.4962	-0.048	0.496214
X Variable 2	0.05192	0.101297	0.5126	0.61938	-0.1738	0.2776	-0.174	0.277624
X Variable 3	1.21646	0.702718	1.7311	0.11412	-0.3493	2.7822	-0.349	2.782211
X Variable 4	0.186	0.78296	0.2376	0.81702	-1.5585	1.9305	-1.559	1.930547

April 1 Snowpack, Linear Regression Based on 13 Historical Seasons

Regression (non-seeded) period:

<u>Water Year</u>	<u>Control avg</u>	<u>Target avg</u>
1975	29.6	9.9
1976	24.8	10.0
1977	10.2	3.6
1978	29.9	10.5
1979	28.6	14.6
1980	35.3	18.4
1981	16.2	9.5
1982	34.9	14.0
1983	31.9	17.0
1984	27.8	12.2
1985	25.0	11.4
1986	35.1	14.3
1987	14.5	10.4
Mean	26.4	12.0

Seeded period:

<u>Water Year</u>	<u>Control Avg</u>	<u>Target Avg</u>	<u>Predicted</u>	<u>Ratio</u>	<u>Increase</u>
1989*	24.5	9.0	11.2	0.80	-2.3
1990*	18.6	10.6	9.0	1.18	1.6
1991*	19.9	10.1	9.5	1.06	0.6
1992*	13.8	8.4	7.2	1.16	1.2
1993*	29.2	14.6	13.0	1.12	1.6
1995*	28.7	15.2	12.8	1.19	2.4
2001*	16.6	10.2	8.3	1.23	1.9
2002*	21.2	6.8	10.0	0.68	-3.2
2003	17.0	9.4	8.4	1.11	1.0
2004**	24.6	7.9	11.3	0.70	-3.4
2005	37.0	20.5	15.9	1.29	4.6
2006	35.4	11.0	15.4	0.72	-4.3
2007**	16.7	6.5	8.3	0.79	-1.8
2008	27.4	11.9	12.3	0.97	-0.4
2009	28.5	7.7	12.7	0.60	-5.0
2010	17.2	9.4	8.5	1.11	0.9
2011	41.6	14.1	17.7	0.80	-3.6
2012**	16.1	5.9	8.1	0.73	-2.2
2013	17.4	7.0	8.6	0.81	-1.6
2015**	12.6	2.3	6.8	0.34	-4.4
2016	21.7	10.1	10.2	0.99	-0.1
2017	32.0	14.8	14.1	1.05	0.7

2018	14.2	6.9	7.4	0.93	-0.5
2019	30.8	14.3	13.6	1.05	0.6
2020	24.1	12.7	11.1	1.15	1.6
2021	17.8	8.6	8.7	0.99	-0.1
2022	15.2	9.9	7.7	1.28	2.1
Seeded Mean	25.0	11.0	11.4	0.96	-0.5

* Seeding conducted in nearby areas but not in target area

** Not included in average due to very early and abnormal snow melt

SUMMARY OUTPUT

<i>Regression Statistics</i>	
Multiple R	0.807491
R Square	0.652042
Adjusted R Square	0.62041
Standard Error	2.344172
Observations	13

	<i>Coefficients</i>	<i>Standard Error</i>	<i>t Stat</i>	<i>P-value</i>	<i>Lower 95%</i>
Intercept	2.028078	2.285175	0.887493	0.393805	-3.00156
X Variable 1	0.376232	0.082868	4.540157	0.000844	0.193842

April 1 Snowpack, Multiple Linear Regression Based on 13 Historical Seasons

Regression (non-seeded)

period:

Water

<u>Year</u>	<u>Timp Div</u>	<u>Farm Cyn</u>	<u>Lookout</u>	<u>Kelley RS</u>	<u>Target Avg</u>
1975	31.6	40.6	25.8	20.5	9.9
1976	26.5	34.2	19.0	19.3	10.0
1977	7.9	17.6	8.8	6.5	3.6
1978	32.3	38.8	24.1	24.4	10.5
1979	33.2	38.7	24.8	17.7	14.6
1980	40.5	43.4	35.1	22.2	18.4
1981	18.3	24.0	13.5	8.9	9.5
1982	39.2	44.1	32.8	23.4	14.0
1983	36.6	43.5	29.9	17.6	17.0
1984	27.0	38.3	26.8	19.0	12.2
1985	25.1	34.3	26.7	13.9	11.4
1986	39.6	43.0	30.2	27.6	14.3
1987	11.6	20.1	16.9	9.3	10.4

Mean	28.4	35.4	24.2	17.7	12.0
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Seeded period:

<u>Water Year</u>	<u>Timp Div</u>	<u>Farm Cyn</u>	<u>Lookout</u>	<u>Kelley RS</u>	<u>Target Avg</u>	<u>Predicted</u>	<u>Ratio</u>	<u>Increase</u>
1989*	19.3	36.5	25.3	16.8	9.0	7.4	1.21	1.5
1990*	21.7	23.7	16.4	12.4	10.6	11.5	0.93	-0.8
1991*	18.3	28.6	20.4	12.4	10.1	9.4	1.08	0.7
1992*	10.1	21.1	12.9	11.0	8.4	5.3	1.58	3.1
1993*	37.1	35.1	27.0	17.7	14.6	17.9	0.82	-3.2
1995*	28.0	39.2	31.5	15.9	15.2	13.8	1.10	1.4
2001*	8.2	27.5	20.3	10.5	10.2	5.0	2.05	5.2
2002*	13.9	34.0	24.1	12.7	6.8	6.4	1.07	0.5
2003	10.7	23.2	20.3	13.8	9.4	6.8	1.39	2.6
2004**	16.7	40.9	28.2	12.7	7.9	6.9	1.14	1.0
2005	40.6	53.1	36.6	17.5	20.5	16.9	1.21	3.6
2006	26.3	53.2	41.7	20.5	11.0	10.2	1.08	0.8
2007**	10.3	24.0	19.4	13.0	6.5	6.2	1.05	0.3
2008	26.7	37.7	29.5	15.6	11.9	13.0	0.92	-1.1
2009	23.6	43.8	30.3	16.3	7.7	9.2	0.84	-1.5
2010	17.8	22.9	18.2	9.8	9.4	11.2	0.84	-1.8
2011	43.7	56.4	44.6	21.5	14.1	19.1	0.73	-5.1
2012**	12.9	20.8	17.8	12.7	5.9	8.2	0.71	-2.4
2014	12.7	31.7	28.2	19.1	7.5	6.0	1.25	1.5
2015**	4.8	20.0	14.1	11.5	2.3	3.2	0.74	-0.8
2016	16.5	30.4	25.4	14.2	10.1	9.1	1.11	1.0
2017	29.2	39.8	33.9	25.0	14.8	12.1	1.22	2.7
2018	8.8	19.6	15.2	13.2	6.9	5.3	1.29	1.6
2019	32.5	41.0	35.0	14.8	14.3	17.3	0.82	-3.0
2020	17.9	31.7	31.1	15.8	12.7	11.0	1.15	1.7
2021	12.7	22.5	23.8	12.0	8.6	10.1	0.86	-1.5
2022	12.7	21.1	17.2	9.8	9.9	8.8	1.13	1.1

Seeded

Mean	21.4	34.9	28.1	15.6	11.0	10.6	1.04	0.4
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* Seeding conducted in nearby areas but not in target area

** Not included in average due to very early and abnormal snow melt

SUMMARY OUTPUT

<i>Regression Statistics</i>	
Multiple R	0.93716
R Square	0.878269
Adjusted R Square	0.817404
Standard Error	1.625839
Observations	13

					Lower Upper 95.0 95.0%			
	Coefficients	Standard Error	t Stat	P-value	Lower 95%	Upper 95%	95.0 %	Upper 95.0%
Intercept	6.339979	3.026456	2.094853	0.069492	0.63905	- 13.31	10.63	13.319
Timp Div	0.536956	0.221169	2.427815	0.041343	0.02694	1.046	0.02	1.046972
Farm Cyn	-0.36777	0.264512	-1.39037	0.201875	0.97774	- 0.242	0.97	0.242197
Lookout	0.388727	0.169898	2.288	0.051425	0.00306	- 0.780	0.00	0.780512
Kelley RS	-0.33837	0.174272	-1.9416	0.088128	0.74024	- 0.063	0.74	0.063505

April 1 Snowpack, Linear Regression Based on 46 Historical Seasons

Regression (non-seeded) period:

<u>Water Year</u>	<u>Control Avg</u>	<u>Target Avg</u>
1957	25.85	7.97
1958	32.65	10.80
1959	18.20	7.90
1960	22.35	5.87
1961	16.30	5.20
1962	32.75	16.23
1963	17.80	5.67
1964	20.40	5.27
1965	32.60	9.73
1966	21.75	9.10
1967	27.10	10.23
1968	27.70	10.60
1969	40.05	16.80
1970	24.15	8.07
1971	28.10	9.53
1972	28.25	7.60
1973	31.35	10.90

1974	24.40	5.03
1975	36.10	9.07
1976	30.35	8.93
1977	12.75	2.47
1978	35.55	9.87
1979	35.95	13.03
1980	41.95	17.67
1981	21.15	8.03
1982	41.65	12.50
1983	40.05	16.40
1984	32.65	11.50
1985	29.70	10.40
1986	41.30	12.53
1987	15.85	7.40
1988	13.40	5.27
1989	27.90	7.27
1990	22.70	8.60
1991	23.45	9.37
1992	15.60	7.07
1993	36.10	14.07
1994	21.90	7.70
1996	28.05	8.03
1997	43.90	13.50
1998	33.35	10.10
1999	21.35	6.00
2000	28.60	10.33
2001	17.85	8.63
2002	23.95	5.93

Mean	27.8	9.5
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Seeded period:

<u>Water Year</u>	<u>Control Avg</u>	<u>Target Avg</u>	<u>Predicted</u>	<u>Ratio</u>	<u>Increase</u>
2003	16.95	8.93	5.78	1.55	3.2
2004**	28.80	6.30	9.86	0.64	-3.6
2005	46.85	19.63	16.09	1.22	3.5
2006	39.75	10.33	13.64	0.76	-3.3
2007**	17.15	3.93	5.84	0.67	-1.9
2008	32.20	11.70	11.04	1.06	0.7
2009	33.70	6.67	11.55	0.58	-4.9
2010	20.35	8.07	6.95	1.16	1.1
2011	50.05	13.57	17.19	0.79	-3.6
2012**	16.85	3.87	5.74	0.67	-1.9
2013	20.20	6.10	6.90	0.88	-0.8
2014	22.20	6.47	7.59	0.85	-1.1
2015**	12.40	1.50	4.21	0.36	-2.7
2016	23.45	8.60	8.02	1.07	0.6
2017	34.50	13.77	11.83	1.16	1.9

2018	14.20	4.83	4.83	1.00	0.0
2019	36.75	13.57	12.60	1.08	1.0
2020	24.80	11.07	8.48	1.30	2.6
2021	17.60	7.13	6.00	1.19	1.1
2022	16.90	9.10	5.76	1.58	3.3

Seeded					
Mean	28.2	10.0	9.6	1.03	0.3

** Not included in average due to very early snow melt
SUMMARY OUTPUT

<i>Regression Statistics</i>			
Multiple R	0.836371208		
R Square	0.699516797		
Adjusted R Square	0.692687634		
Standard Error	1.885329949		
Observations	46		

	<i>Coefficients</i>	<i>Standard Error</i>
Intercept	-0.07114187	0.987139943
X Variable 1	0.344927472	0.034081011

April 1 Snowpack, Multiple Linear Regression Based on 46 Historical Seasons

<u>Water</u>			
<u>Year</u>	<u>Farmington Cyn</u>	<u>Timpanogos</u>	<u>Target Avg</u>
1957	26.40	25.30	7.97
1958	33.90	31.40	10.80
1959	21.10	15.30	7.90
1960	25.40	19.30	5.87
1961	21.80	10.80	5.20
1962	35.40	30.10	16.23
1963	20.50	15.10	5.67
1964	23.90	16.90	5.27
1965	38.60	26.60	9.73
1966	22.10	21.40	9.10
1967	23.00	31.20	10.23
1968	30.50	24.90	10.60
1969	36.40	43.70	16.80
1970	30.30	18.00	8.07

1971	38.70	17.50	9.53
1972	37.60	18.90	7.60
1973	33.70	29.00	10.90
1974	30.90	17.90	5.03
1975	40.60	31.60	9.07
1976	34.20	26.50	8.93
1977	17.60	7.90	2.47
1978	38.80	32.30	9.87
1979	38.70	33.20	13.03
1980	43.40	40.50	17.67
1981	24.00	18.30	8.03
1982	44.10	39.20	12.50
1983	43.50	36.60	16.40
1984	38.30	27.00	11.50
1985	34.30	25.10	10.40
1986	43.00	39.60	12.53
1987	20.10	11.60	7.40
1988	16.10	10.70	5.27
1989	36.50	19.30	7.27
1990	23.70	21.70	8.60
1991	28.60	18.30	9.37
1992	21.10	10.10	7.07
1993	35.10	37.10	14.07
1994	25.70	18.10	7.70
1995	39.20	28.00	13.53
1997	51.60	36.20	13.50
1998	43.50	23.20	10.10
1999	27.50	15.20	6.00
2000	39.20	18.00	10.33
2001	27.50	8.20	8.63
2002	34.00	13.90	5.93

Mean	32.0	23.5	9.5
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Water

<u>Year</u>	<u>Farmington Cyn</u>	<u>Timpanogos</u>	<u>Target avg</u>	<u>Predicted</u>	<u>Ratio</u>	<u>Increase</u>
2003	23.20	10.70	8.93	5.51	1.62	3.4
2004**	40.90	16.70	6.30	8.28	0.76	-2.0
2005	53.10	40.60	19.63	15.45	1.27	4.2
2006	53.20	26.30	10.33	11.65	0.89	-1.3
2007**	24.00	10.30	3.93	5.46	0.72	-1.5
2008	37.70	26.70	11.70	10.73	1.09	1.0
2009	43.80	23.60	6.67	10.31	0.65	-3.6
2010	22.90	17.80	8.07	7.38	1.09	0.7
2011	56.40	43.70	13.57	16.49	0.82	-2.9
2012**	20.80	12.90	3.87	5.94	0.65	-2.1
2013	30.70	9.70	6.10	5.74	1.06	0.4
2014	31.70	12.70	6.47	6.61	0.98	-0.1

2015**	20.00	4.80	1.50	3.73	0.40	-2.2
2016	30.40	16.50	8.60	7.53	1.14	1.1
2017	39.80	29.20	13.77	11.53	1.19	2.2
2018	19.60	8.80	4.83	4.77	1.01	0.1
2019	41.00	32.50	13.57	12.49	1.09	1.1
2020	31.70	17.90	11.07	7.99	1.38	3.1
2021	22.50	12.70	7.13	6.00	1.19	1.1
2022	21.10	12.70	9.10	5.91	1.54	3.2
Seeded Mean	34.9	21.4	10.0	9.1	1.09	0.8

SUMMARY OUTPUT

<i>Regression Statistics</i>	
Multiple R	0.859094313
R Square	0.738043039
Adjusted R Square	0.719331828
Standard Error	1.801747514
Observations	46

	<i>Coefficients</i>	<i>Standard Error</i>
Intercept	1.132749671	1.060535368
Farmington Cyn	0.066018713	0.045897051
Timpanogos	0.266315504	0.046335651

March – July Streamflow Linear Regression, with 5 Control and 3 Target Sites; units are in acre feet

Regression (non-seeded) Period:

<u>Water Year</u>	<u>Control avg</u>	<u>Target Avg</u>
1966	112936	49949
1971	261215	66992
1972	178150	59875
1973	193597	72462
1974	212877	43409
1975	197588	79701
1976	169736	48415
1977	44359	25649
1978	227917	53303
1979	191656	45339
1983	279948	96463

1984	331384	69498
1985	222233	57727
1986	276152	96943
1987	116536	64515
1988	139135	36566
1989	105895	32889
1990	89112	51965
1991	120377	54937
1992	81594	38662
1993	212713	78967
1994	83576	38992
1995	245111	105683
1996	189341	52819
1997	263786	76363
1998	215275	81533
1999	215124	75497
2000	120952	40342
2001	113842	62042
2002	58672	19379
Mean	175693	59229

Seeded Period:

<u>Water Year</u>	<u>Control Avg</u>	<u>Target Avg</u>	<u>Predicted</u>	<u>Ratio</u>	<u>Increase</u>
2003	123438	47931	47895	1.00	36
2004	90888	40375	40836	0.99	-460
2005	174888	101668	59055	1.72	42614
2006	152841	54263	54273	1.00	-10
2007	105346	33724	43971	0.77	-10248
2008	207348	45549	66095	0.69	-20546
2009	219964	54665	68831	0.79	-14166
2010	175017	51930	59082	0.88	-7152
2011	365025	103727	100293	1.03	3433
2012	79824	29931	38436	0.78	-8505
2013	80584	36523	38601	0.95	-2077
2014	177875	35639	59702	0.60	-24063
2015	149671	51525	53585	0.96	-2060
2016	178270	61738	59788	1.03	1950
2017	189133	83172	62144	1.34	21028
2018	94881	30575	41702	0.73	-11127
2019	192441	85429	62861	1.36	22567
2020	128875	51130	49075	1.04	2056
2021	60566	25790	34259	0.75	-8469

Seeded Mean	155099	53962	54762	0.99	-800
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SUMMARY OUTPUT

<i>Regression Statistics</i>	
Multiple R	0.749069335
R Square	0.561104868
Adjusted R Square	0.545430042
Standard Error	14338.66364
Observations	30

	<i>Coefficients</i>	<i>Standard Error</i>	<i>t Stat</i>	<i>P-value</i>	<i>Lower 95%</i>
Intercept	21123.14391	6886.056	3.06752417	0.00475	7017.681
X Variable 1	0.216890053	0.036251	5.98302272	1.92E-06	0.142633

**High Uintas March – July Streamflow Multiple Linear Regression, with 5 Control and 3 Target Sites;
units are in acre feet**

Regression (non-seeded) Period:

<u>Water Year</u>	<u>Hams Fk</u>	<u>Fonten elle</u>	<u>Smiths Fk</u>	<u>Little Snake</u>	<u>White River</u>	<u>Target Avg</u>
1966	44794	26481	69071	261819	162515	49949
1971	145432	70383	178721	590894	320645	66992
1972	103820	75862	158637	301634	250798	59875
1973	48082	29485	75594	476355	338467	72462
1974	80404	46964	127332	498440	311243	43409
1975	81706	45447	115301	396510	348975	79701
1976	75548	52151	120425	329104	271451	48415
1977	7077	7711	23732	85711	97566	25649
1978	93460	58383	142896	471055	373789	53303
1979	53667	33706	80654	396038	394214	45339
1983	102494	73684	153030	617221	453311	96463
1984	103004	56974	147686	809511	539744	69498
1985	49380	32445	86070	470868	472404	57727
1986	128700	95836	186880	499949	469394	96943
1987	36867	24696	51531	219782	249806	64515
1988	36184	24103	64874	298988	271525	36566
1989	46081	30952	84247	170223	197970	32889
1990	33395	23630	62426	171219	154892	51965
1991	44451	23899	77260	213547	242727	54937

1992	23469	10950	48549	140134	184870	38662
1993	69422	33656	122948	457750	379790	78967
1994	27123	17019	46243	176877	150618	38992
1995	57851	40953	106167	564912	455670	105683
1996	72113	40088	129123	364185	341195	52819
1997	91551	59499	165808	589422	412650	76363
1998	58520	41232	102936	458203	415485	81533
1999	80859	69012	137185	480812	307753	75497
2000	37484	23018	70236	244056	229966	40342
2001	20646	14235	44049	238488	251794	62042
2002	24183	18504	49405	93630	107637	19379

Seeded

Mean	62592	40032	100967	369578	305295	59229
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Seeded Period:

<u>Water Year</u>	<u>Hams Fork</u>	<u>Fontene lle</u>	<u>Smiths Fork</u>	<u>Little Snake</u>	<u>White River</u>	<u>Target Avg</u>	<u>Predicted</u>	<u>Ratio</u>	<u>Increase</u>
2003	26134	29925	57863	242638	260630	47931	57296	0.84	-9365
2004	30335	23304	60098	152754	187948	40375	43108	0.94	-2733
2005	57851	53163	113152	322611	309446	101668	66848	1.52	34820
2006	72113	43893	95628	235021	332619	54263	59780	0.91	-5516
2007	91551	19643	52585	215647	209043	33724	32320	1.04	1404
2008	58520	33729	81623	512575	353108	45549	65572	0.69	-20023
2009	80859	41152	117741	542915	332130	54665	61026	0.90	-6361
2010	37484	34226	71247	470661	251381	51930	60795	0.85	-8865
2011	20646	82651	159392	943100	534183	103727	124156	0.84	-20430
2012	24183	23792	64335	134015	138679	29931	39259	0.76	-9328
2013	26134	17708	57232	121059	180197	36523	39173	0.93	-2650
2014	81110	53750	107247	324809	322459	35639	64300	0.55	-28661
2015	58245	39208	95950	237787	317166	51525	58174	0.89	-6649
2016	58245	34884	90428	420466	286091	61738	57181	1.08	4557
2017	58245	93600	183955	307629	237851	83172	79093	1.05	4079
2018	58245	32380	90296	142410	162616	30575	37738	0.81	-7163
2019	58245	35265	91752	394096	396364	85429	66959	1.28	18469
2020	58245	31271	83105	292991	196115	51130	44272	1.15	6858
2021	58245	16921	47579	98823	117765	25790	26708	0.97	-918
Avg	55511	38972	90590	321684	269779	53962	56608	0.95	-2645

SUMMARY OUTPUT

<i>Regression Statistics</i>	
Multiple R	0.788019316
R Square	0.620974442
Adjusted R Square	0.542010784
Standard Error	14392.49006
Observations	30
<i>Coefficients</i>	
Intercept	19093.5744
Hams Fork	-0.20489592
Fontenelle	0.648935056
Smiths Fork	-0.09760667
Little Snake	0.022804631
White River	0.093055464

High Uintas March – July Streamflow Linear Regression, with 3 Wyoming Control Sites and 3 Target Sites; units are in acre feet

Regression (non-seeded) Period:

<u>Water Year</u>	<u>Control Avg</u>	<u>Target Avg</u>
1966	46782	49949
1971	131512	66992
1972	112773	59875
1973	51054	72462
1974	84900	43409
1975	80818	79701
1976	82708	48415
1977	12840	25649
1978	98246	53303
1979	56009	45339
1983	109736	96463
1984	102555	69498
1985	55965	57727
1986	137139	96943
1987	37698	64515
1988	41720	36566
1989	53760	32889
1990	39817	51965
1991	48537	54937
1992	27656	38662

1993	75342	78967
1994	30128	38992
1995	68324	105683
1996	80441	52819
1997	105619	76363
1998	67563	81533
1999	95685	75497
2000	43579	40342
2001	26310	62042
2002	30697	19379

<u>Mean</u>	<u>67864</u>	<u>59229</u>
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Seeded Period:

<u>Water Year</u>	<u>Control Avg</u>	<u>Target Avg</u>	<u>Predicted</u>	<u>Ratio</u>	<u>Increase</u>
2003	37974	47931	47492	1.01	439
2004	37912	40375	47468	0.85	-7093
2005	80795	101668	64307	1.58	37361
2006	65521	54263	58309	0.93	-4046
2007	34013	33724	45937	0.73	-12213
2008	57019	45549	54971	0.83	-9422
2009	74925	54665	62002	0.88	-7337
2010	51014	51930	52613	0.99	-683
2011	115947	103727	78110	1.33	25616
2012	42142	29931	49129	0.61	-19198
2013	33887	36523	45888	0.80	-9364
2014	80702	35639	64271	0.55	-28632
2015	64468	51525	57896	0.89	-6370
2016	61598	61738	56769	1.09	4969
2017	133395	83172	84962	0.98	-1790
2018	56460	30575	54751	0.56	-24177
2019	57248	85429	55061	1.55	30368
2020	51756	51130	52904	0.97	-1774
2021	28747	25790	43869	0.59	-18079
Seeded Mean	61343	53962	56669	0.95	-2706

SUMMARY OUTPUT

<u>Regression Statistics</u>	
Multiple R	0.609172
R Square	0.371091
Adjusted R Square	0.34863
Standard Error	17164.15

Observations 30

	<i>Coefficients</i>	<i>Standard Error</i>	<i>t Stat</i>	<i>P-value</i>	<i>Lower 95%</i>
Intercept	32580.86	7266.534	4.48368633	0.000114	17696.02
X Variable 1	0.392674	0.096607	4.06467088	0.000353	0.194784

High Uintas March – July Streamflow Multiple Linear Regression, with 3 Wyoming Control Sites and 3 Target Sites; units are in acre feet

Regression (non-seeded) Period:

<u>Water</u>	<u>Hams</u>			
<u>Year</u>	<u>Fork</u>	<u>Fontenelle</u>	<u>Smiths Fork</u>	<u>Target Avg</u>
1966	44794	26481	69071	49949
1971	145432	70383	178721	66992
1972	103820	75862	158637	59875
1973	48082	29485	75594	72462
1974	80404	46964	127332	43409
1975	81706	45447	115301	79701
1976	75548	52151	120425	48415
1977	7077	7711	23732	25649
1978	93460	58383	142896	53303
1979	53667	33706	80654	45339
1983	102494	73684	153030	96463
1984	103004	56974	147686	69498
1985	49380	32445	86070	57727
1986	128700	95836	186880	96943
1987	36867	24696	51531	64515
1988	36184	24103	64874	36566
1989	46081	30952	84247	32889
1990	33395	23630	62426	51965
1991	44451	23899	77260	54937
1992	23469	10950	48549	38662
1993	69422	33656	122948	78967
1994	27123	17019	46243	38992
1995	57851	40953	106167	105683
1996	72113	40088	129123	52819
1997	91551	59499	165808	76363
1998	58520	41232	102936	81533
1999	80859	69012	137185	75497
2000	37484	23018	70236	40342
2001	20646	14235	44049	62042
2002	24183	18504	49405	19379
Average	62592	40032	100967	59229

Seeded Period:

<u>Water</u>	<u>Hams</u>						
<u>Year</u>	<u>Fork</u>	<u>Fontenelle</u>	<u>Smiths Fork</u>	<u>Target Avg</u>	<u>Predicted</u>	<u>Ratio</u>	<u>Increase</u>
2003	26134	29925	57863	47931	52856	0.91	-4924
2004	30335	23304	60098	40375	49158	0.82	-8782
2005	76070	53163	113152	101668	65107	1.56	36561
2006	57043	43893	95628	54263	61159	0.89	-6896
2007	29811	19643	52585	33724	45597	0.74	-11874
2008	55706	33729	81623	45549	53019	0.86	-7470
2009	65884	41152	117741	54665	63245	0.86	-8580
2010	47569	34226	71247	51930	52717	0.99	-787
2011	105799	82651	159392	103727	83343	1.24	20384
2012	38298	23792	64335	29931	48385	0.62	-18453
2013	26722	17708	57232	36523	46693	0.78	-10170
2014	81110	53750	107247	35639	62524	0.57	-26885
2015	58245	39208	95950	51525	58683	0.88	-7158
2016	59483	34884	90428	61738	54856	1.13	6882
2017	122630	93600	183955	83172	90500	0.92	-7327
2018	46705	32380	90296	30575	57002	0.54	-26427
2019	44728	35265	91752	85429	59282	1.44	26147
2020	40893	31271	83105	51130	56143	0.91	-5013
2021	21740	16921	47579	25790	45130	0.57	-19340
Seeded							
Mean	54469	38972	90590	53962	58179	0.93	-4216

SUMMARY OUTPUT

<i>Regression Statistics</i>	
Multiple R	0.629376912
R Square	0.396115297
Adjusted R Square	0.326436293
Standard Error	17454.10769
Observations	30

	<i>Coefficients</i>	<i>Standard Error</i>
Intercept	30446.25283	9346.848154
Hams Fork	-0.26435458	0.430607215
Fontenelle	0.478306208	0.486930199
Smiths Fork	0.259311656	0.340472822

APPENDIX E: GLOSSARY OF RELEVANT METEOROLOGICAL TERMS

Advection: Movement of an air mass. Cold advection describes a colder air mass moving into the area, and warm advection is used to describe an incoming warmer air mass. Dry and moist advection can be used similarly.

Air Mass: A term used to describe a region of the atmosphere with certain defining characteristics. For example, a cold or warm air mass, or a wet or dry air mass. It is a fairly subjective term but is usually used in reference to large (synoptic scale) regions of the atmosphere, both near the surface and/or at mid and upper levels of the atmosphere.

Cold-core low: A typical mid-latitude type of low pressure system, where the core of the system is colder than its surroundings. This type of system is also defined by the cyclonic circulation being strongest in the upper levels of the atmosphere. The opposite is a warm-core low, which typically occurs in the tropics.

Cold Pool: An air mass that is cold relative to its surroundings, and may be confined to a particular basin

Condensation: Phase change of water vapor into liquid form. This can occur on the surface of objects (such as dew on the grass) or in mid-air (leading to the formation of clouds). Clouds are technically composed of water in liquid form, not water vapor.

Confluent: Wind vectors coming closer together in a two-dimensional frame of reference (opposite of diffluent). The term convergence is also used similarly.

Convective (or convection): Pertains to the development of precipitation areas due to the rising of warmer, moist air through the surrounding air mass. The warmth and moisture contained in a given air mass makes it lighter than colder, dryer air. Convection often leads to small-scale, locally heavy showers or thundershowers. The opposite precipitation type is known as stratiform precipitation.

Convergence: Refers to the converging of wind vectors at a given level of the atmosphere. Low-level convergence (along with upper-level divergence), for instance, is associated with lifting of the air mass which usually leads to development of clouds and precipitation. Low-level divergence (and upper-level convergence) is associated with atmospheric subsidence, which leads to drying and warming.

Deposition: A phase change where water vapor turns directly to solid form (ice). The opposite process is called sublimation.

Dew point: The temperature at which condensation occurs (or would occur) with a given amount of moisture in the air.

Diffluent: Wind vectors spreading further apart in a two-dimensional frame of reference; opposite of confluent

Entrain: Usually used in reference to the process of a given air mass being ingested into a storm system

Evaporation: Phase change of liquid water into water vapor. Water vapor is usually invisible to the eye.

El Nino: A reference to a particular phase of oceanic and atmospheric temperature and circulation patterns in the tropical Pacific, where the prevailing easterly trade winds weaken or dissipate. Often has an effect on mid-latitude patterns as well, such as increased precipitation in southern portions of the U.S. and decreased precipitation further north. The opposite phase is called La Nina.

Front (or frontal zone): Reference to a temperature boundary with either incoming colder air (cold front) or incoming warmer air (warm front); can sometimes be a reference to a stationary temperature boundary line (stationary front) or a more complex type known as an occluded front (where the temperature change across a boundary can vary in type at different elevations).

Glaciogenic: Ice-forming (aiding the process of nucleation); usually used in reference to cloud seeding nuclei

GMT (or UTC, or Z) time: Greenwich Mean Time, universal time zone corresponding to the time at Greenwich, England. Pacific Standard Time (PST) = GMT – 8 hours; Pacific Daylight Time (PDT) = GMT – 7 hours.

Graupel: A precipitation type that can be described as “soft hail”, that develops due to riming (nucleation around a central core). It is composed of opaque (white) ice, not clear hard ice such as that contained in hailstones. It usually indicated the presence of convective clouds and can be associated with electrical charge separation and occasionally lightning activity.

High Pressure (or Ridge): Region of the atmosphere usually accompanied by dry and stable weather. Corresponds to a northward bulge of the jet stream on a weather map, and to an anti-cyclonic (clockwise) circulation pattern.

Inversion: Refers to a layer of the atmosphere in which the temperature increase with elevation

Jet Stream or Upper-Level Jet (sometimes referred to more generally as the storm track): A region of maximum wind speed, usually in the upper atmosphere that usually coincides with the main storm track in the mid-latitudes. This is the area that also typically corresponds to the greatest amount of mid-latitude synoptic-scale storm development.

La Nina: The opposite phase of that known as El Nino in the tropical Pacific. During La Nina the easterly tropical trade winds strengthen and can lead in turn to a strong mid-latitude storm track, which often brings wetter weather to northern portions of the U.S.

Longwave (or longwave pattern): The longer wavelengths, typically on the order of 1,000 – 2,000+ miles of the typical ridge/trough pattern around the northern (or southern) Hemisphere, typically most pronounced in the mid-latitudes.

Low-Level Jet: A zone of maximum wind speed in the lower atmosphere. Can be caused by geographical features or various weather patterns, and can influence storm behavior and dispersion of cloud seeding materials

Low-pressure (or trough): Region of the atmosphere usually associated with stormy weather. Corresponds to a southward dip to the jet stream on a weather map as well as a cyclonic (counter-clockwise) circulation pattern in the Northern Hemisphere.

Mesoscale: Sub - synoptic scale, about 100 miles or less; this is the size scale of more localized weather features (such as thunderstorms or mountain-induced weather processes).

Microphysics: Used in reference to composition and particle types in a cloud

MSL (Mean Sea Level): Elevation height reference in comparison to sea level

Negative (ly) tilted trough: A low-pressure trough where a portion is undercut, such that a frontal zone can be in a northwest to southeast orientation.

Nucleation: The process of supercooled water droplets in a cloud turning to ice. This is the process that is aided by cloud seeding. For purposes of cloud seeding, there are three possible types of cloud composition: Liquid (temperature above the freezing point), supercooled (below freezing but still in liquid form), and ice crystals.

Nuclei: Small particles that aid water droplet or ice particle formation in a cloud

Orographic: Terrain-induced weather processes, such as cloud or precipitation development on the upwind side of a mountain range. Orographic lift refers to the lifting of an air mass as it encounters a mountain range.

Pressure Heights:

(700 millibars, or mb): Corresponds to approximately 10,000 feet above sea level (MSL); 850 mb corresponds to about 5,000 feet MSL; and 500 mb corresponds to about 18,000 feet MSL. These are standard height levels that are occasionally referenced, with the 700-mb level most important regarding cloud-seeding potential in most of the western U.S.

Positive (ly) tilted trough: A normal U-shaped trough configuration, where an incoming cold front would generally be in a northeast– southwest orientation.

Reflectivity: The density of returned signal from a radar beam, which is typically bounced back due to interaction with precipitation particles (either frozen or liquid) in the atmosphere. The reflectivity depends on the size, number, and type of particles that the radar beam encounters

Ridge (or High Pressure System): Region of the atmosphere usually accompanied by dry and stable weather. Corresponds to a northward bulge of the jet stream on a weather map, and to an anti-cyclonic (clockwise) circulation pattern.

Ridge axis: The longitude band corresponding to the high point of a ridge

Rime (or rime ice): Ice buildup on an object (often on an existing precipitation particle) due to the freezing of supercooled water droplets.

Shortwave (or shortwave pattern): Smaller-scale wave features of the weather pattern typically seen at mid-latitudes, usually on the order of a few to several hundred miles; these often correspond to individual frontal systems

Silver iodide: A compound commonly used in cloud seeding because of the similarity of its molecular structure to that of an ice crystal. This structure helps in the process of nucleation, where supercooled cloud water changes to ice crystal form.

Storm Track (sometimes reference as the Jet Stream): A zone of maximum storm propagation and development, usually concentrated in the mid-latitudes.

Stratiform: Usually used in reference to precipitation, this implies a large area of precipitation that has a fairly uniform intensity except where influenced by terrain, etc. It is the result of larger-scale (synoptic scale) weather processes, as opposed to convective processes.

Sublimation: The phase change in which water in solid form (ice) turns directly into water vapor. The opposite process is deposition.

Subsidence: The process of a given air mass moving downward in elevation, such as often occurs on the downwind side of a mountain range

Supercooled: Liquid water (such as tiny cloud droplets) occurring at temperatures below the freezing point (32 F or 0 C).

Synoptic Scale: A scale of hundreds to perhaps 1,000+ miles, the size scale at which high and low pressure systems develop

Trough (or low pressure system): Region of the atmosphere usually associated with stormy weather. Corresponds to a southward dip to the jet stream on a weather map as well as a cyclonic (counter-clockwise) circulation pattern in the Northern Hemisphere.

Trough axis: The longitude band corresponding to the low point of a trough

Upper-Level Jet or Jet Stream (sometimes referred to more generally as the storm track): A region of maximum wind speed, usually in the upper atmosphere that usually coincides with the main storm track in the mid-latitudes. This is the area that also typically corresponds to the greatest amount of mid-latitude synoptic-scale storm development.

UTC (or GMT, or Z) time: Greenwich Mean Time, universal time zone corresponding to the time at Greenwich, England. Pacific Standard Time (PST) = GMT – 8 hours; Pacific Daylight Time (PDT) = GMT – 7 hours.

Vector: Term used to represent wind velocity (speed + direction) at a given point

Velocity: Describes speed of an object, often used in the description of wind intensities

Vertical Wind Profiler: Ground-based system that measures wind velocity at various levels above the site